

**In situ experiments on nuclide
migration in fractured crystalline
rocks**

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Studsvik Energiteknik and
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IN SITU EXPERIMENTS ON NUCLIDE MIGRATION IN FRACTURED
CRYSTALLINE ROCKS

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1. INTRODUCTION

The migration rate of a nuclide with groundwater in geological media is lowered, as compared to the water movement, by different reactions, e.g. ion exchange, adsorption, precipitation and mineralization. The retardation of a nuclide "i" is usually expressed by the retardation factor K_i , which is defined as the ratio between the water velocity u_p (in pores) and the velocity u_i of the nuclide and related to the mass distribution coefficient K_d ($m^3 \cdot kg^{-1}$) by the equation

$$K_i = \frac{u_p}{u_i} = 1 + K_d \cdot \frac{\rho}{\epsilon} \quad (1)$$

where ρ is the bulk density ($kg \cdot m^{-3}$) of the dry porous medium and ϵ is the porosity. Equation (1) is valid for porous geological media such as sand and unconsolidated sediments, whereas for fractured rocks the following equation is often used:

$$K_i = \frac{u_p}{u_i} = 1 + K_a \cdot R_f \quad (2)$$

where K_a is the surface distribution coefficient ($m^3 \cdot m^{-2}$) and R_f is the surface-to-volume ratio of the fractures.

The retardation factor K_i is usually calculated from K_d and K_a values determined in laboratory experiments with rock or soil samples. Such samples cannot, however, in all respects be representative of the migration media. Migration experiments carried out "in situ" are therefore an important complement to laboratory studies.

In the present paper are reported the results and experience of the first two of a series of field experiments carried out in natural non-disturbed fractures of crystalline rocks. The aim of the first series was to study the migration of some selected radionuclides and to develop suitable methods and techniques for such studies. The aim of the second series was to test the effect of bentonite injection into rock fractures on the radionuclide migration. The experiments were carried out in co-operation between the Geological Survey of Sweden and Studsvik Energiteknik AB.

2. GEOLOGY OF THE TEST AREA

The experimental site is situated close to the Baltic Sea and about 90 km south of Stockholm. The rocks in this area, which are mainly of Precambrian age, are described by Lundström (1976).

The migration experiments were performed in a large outcrop of migmatitic-gneissic granite. A map of the fractures in the rock surface is shown in fig 1 (from Bjurstedt et al, 1975). Also horizontal fractures were observed in the core of a diamond-drilled borehole, marked as K3 in fig 1. Many of the fractures are filled with calcite and chlorite.

3. HYDROLOGICAL TESTS AND SELECTED TRACER FLOW PATHS

Eight boreholes with a diameter of 115 mm were drilled in the outcrop area and different borehole investigations (including pumping tests, water pressure tests, visual TV logging, resistivity and radioactivity measurements, and ^{82}Br tracer tests) were carried out in order to delineate

the hydrogeological conditions and to find flow paths, suitable for the migration experiments. Three boreholes, marked B2, B7 (observation holes) and B8 (injection hole) in figures 1 and 2 were selected for the experiments. Fig 3 shows the results of sectional permeability determinations in these boreholes, reflecting the fractured nature of the rock.

Two flow paths could be simultaneously utilized in the experiments; one between B8 and B2 (the main flow path) and another between B8 and B7, the distance between the boreholes being 51 and 22 metres, respectively (fig 2). When pumping in B2 (at a depth of 65 metres) with a constant rate of 0.1 litre/second, the peak of the ^{82}Br pulse was recorded at B2 about 40 hours after injection in B8 (fig 4). The corresponding time between B8 and B7 was about 10 hours (fig 4). Special tests with ^{82}Br revealed that the water enters borehole B7 via a fracture zone at 70 metres depth and flows upwards the hole, leaving it through a main fracture zone at about 40 metres depth, probably discharging into the Baltic Sea.

In fig 2 is marked the ground water level (related to the mean sea level) measured on 1977-03-22 at stationary conditions with a pumping rate of 0.1 litre/second. The hydraulic gradient is thus 0.05 between B8 and B2 and 0.11 between B8 and B7. The water pumped from the borehole B2 was allowed to discharge into the Baltic Sea. Neither the detailed configuration of the fracture system nor the distribution of fracture widths are, of course, known. The TV logging revealed one fracture apparently about 1 cm wide in the injection zone of B8, but it is not known how far this width may persist.

4. MIGRATION STUDIES

4.1 Elements studied and radionuclides used

The present study comprised the following elements: selenium, technetium, tin, cesium, iodine, neodymium and strontium. Long-lived nuclides of these elements are present in highly radioactive waste (Table I); neodymium being used as a "stand-in" for the elements americium - curium.

In Table II are listed the isotopes used in the experiments, their nuclear characteristics and MPC values. In addition to the waste elements, one of the macroelements in the groundwater, sodium, was represented by the isotope ^{24}Na . In all the different injections ^{82}Br was used as reference tracer and assumed to follow the water flow with negligible retention.

4.2 Injection procedure

The radionuclides were dissolved in 5 - 10 litres of groundwater, taken from borehole B2, and pumped down between two rubber packers enclosing the high permeability zone at about 72 m depth in borehole B8 (see figures 2 and 3). The nuclides were then forced out into the fractures by rinsing with totally 30 litres of groundwater, spread over three successive operations.

Several radionuclides were usually simultaneously injected. The combinations of radionuclides and injected activities used in the three principal field studies are listed in Table III.

4.3 Radioactivity measurements

The gamma activity of pumped water from borehole B2 was continuously recorded with a Ge(Li) detector. This allowed selective measurements of simultaneously injected isotopes, as illustrated in fig 5 which shows a gamma spectrum measured in connection with the studies of ^{82}Br , ^{131}I and $^{99\text{m}}\text{Tc}$. The net peak areas of the 554.3 keV (^{82}Br), 364.5 keV (^{131}I) and 140.4 keV ($^{99\text{m}}\text{Tc}$) gamma-rays were recorded for the concentration analysis.

Water samples were regularly taken from borehole B2 for calibration and check of the field analysis and for laboratory measurements of those nuclides which emit gamma-rays of too low energy to be measured in the field, e. g. ^{131}Cs .

Also ^{85}Sr was mostly analysed in the laboratory because of interference from the 510 keV gamma-ray in the natural background radiation. In borehole B7 water sampling was made at a level just above the water inflow at 70 metres and analysis was carried out in the laboratory.

The laboratory measurements were mainly based on gamma spectrometry, using a Ge(Li) detector. In the case of very low activity the samples were concentrated by evaporation and measured in a "4 π " geometry using a well-type Ge(Li) detector. The ^{131}Cs activity was measured with a Si(Li) detector after preconcentration and precipitation.

4.4 Results of the migration experiments

In figures 6 - 11 are shown curves of the measured concentration versus migration time for the different nuclides. The concentrations are related to C_0 , which is the tracer concentration in the injected solution. The 22 m flow path is denoted "B8 - B7" and the 51 m flow path "B8 - B2". The measured activities are corrected for radioactive decay and are thus valid for the time of injection.

4.5 Complementary laboratory experiments

Batch experiments and column tests were performed for planning of the field experiments. Complementary laboratory experiments, briefly described in the sequel, were also carried out as an aid to the interpretation of the field data. The results are used in the discussion of the field data later on.

Dried drill cuttings from boreholes B8 and B7 were used, having the following grain size distribution:

Fraction 0.04 - 0.17	81 weight per cent
0.02 - 0.04	12 "-
0.01 - 0.02	5 "-
0.001- 0.01	1 "-

The groundwater used in the experiments was taken from borehole B2 and had the following main composition:

Ca	36	mg/litre	HCO ₃	240	mg/litre
Mg	10	"	SiO ₂	14	"
Na	45	"	pH	7.45	
K	5.5	"	Conducti-	385	μS/cm
			vity		
Cl	11	"			
SO ₄	14	"			

Aliquots of the solutions injected on 1977-03-11 (Table III) were diluted as follows:

A	1 part with 9 parts of groundwater
B	1 part with 4 parts of groundwater and 5 parts of deionized water
C	1 part with 9 parts of deionized water

Batch experiments were performed by shaking different amounts of drill cuttings (15, 50, 150 and 500 g/litre) with solutions A, B and C. After a contact time of two hours the slurries were centrifuged and the activity of the centrifugate measured for K_d calculations. Separate samples of solutions A, B and C were also centrifuged in order to determine the fractions of added activities which were still in solution. The following observations were made:

- Tin is completely hydrolyzed and no K_d could therefore be determined for Sn⁴⁺.
- Most of selenium was precipitated before coming into contact with the cuttings. The soluble selenium was only slightly sorbed on the grains with a K_d of the order of 0 - 3 ml/g.
- Neodymium precipitates to some extent when mixed with the groundwater. The K_d values are thus somewhat uncertain.
- Strontium did not precipitate. A K_d value of 1.5 ml/g was obtained (mean of 5 determinations).

Additional K_d determinations (for I^- , SeO_3^{2-} and Cs^+) were made using groundwater without extra carriers. The results are listed in Table IV together with values obtained in the experiments described above.

4.6 Comments on the results of the field migration experiments

Technetium

Owing to the short half-life of ^{99m}Tc , only the beginning of the breakthrough curve for the flow path B8 - B2 could be recorded, whereas the whole concentration curve was recorded for the shorter flow path B8 - B2, fig 6. As expected, the migration rate for technetium injected as TcO_4^- is the same as that for bromine, i. e. technetium is transported without any observable retention. This is contrary to the results obtained in Oak Ridge for various American rock samples; an appreciable retardation of technetium was observed which is believed to be due to reduction of initially deposited TcO_4^- ions (Carleson, 1978).

Iodine

In fig 7 the breakthrough curve of I^- is compared to that of Br^- for the flow path B8 - B2. The slightly lower maximum of the iodine curve indicates a weak sorption of iodine. Similar results were obtained for the flow path B8 - B7. This indicates a slight difference between the field and the laboratory results ($K_d = 0$, Table IV).

Sodium

The sodium concentration curve for the flow path B8 - B2, shown in fig 8, is very much the same as that of iodine, i. e. a somewhat lower maximum and longer tail as compared to the bromine curve. The sorption process is probably ion exchange which, however, should be weak in view of the relatively high concentrations of calcium and magnesium in the water.

Strontium

As seen from figures 9 and 10, strontium was retarded by a factor of about 6 in both flow paths as compared to bromine. This compares well with the retardation factors of 6 and 3, respectively, which were calculated by Neretnieks (1977) from K_d values determined on drill-cuttings in two laboratories: Allard (1977) and Table IV in this report. The K_d values were recalculated to K_a values assuming spherical grains. Neretnieks obtained fissure data by extending the fracture model of Snow (1968) to assume a normal frequency distribution for the channel width. Fitting the experimental and computed residence time distributions gave a mean fissure width of 0.063 mm with a standard deviation of 0.021 mm for the flow path B8 - B7.

Tin

Applying the results of our laboratory experiments, Sn^{4+} might be expected to partly or completely precipitate in the borehole by hydrolysis. A K_d value of 250 ml/g was reported for desert soil (Burkholder et al 1976). With the exception of some very low activity of ^{113}Sn during the first seven hours in borehole B7, no tin activity was detected until about 20 days later when the beginning of a breakthrough was observed simultaneously with the arrival of selenium activity (fig 11). This seems to indicate either that the tin precipitate, which might be a selenium compound, has a certain solubility or that tin migrates in a chemical form, other than Sn^{4+} , e. g. in colloidal form.

Selenium

Selenium as SeO_3^{2-} is known to form precipitates with Sn^{4+} and under certain conditions also with Ca^{2+} , which complicates the interpretation. Selenium was injected partly as precipitates and partly in solution as SeO_3^{2-} . In borehole B7 as well as in B2 selenium arrived simultaneously with bromine (fig 11). The concentration decreased rapidly to almost zero and increased again after about 20 days in a way similar

to that of tin. This indicates different migration rates of selenium depending on its chemical form. The fact that selenium and tin arrive close together in their second pulse suggests their migration in a similar chemical form. The batch experiments gave values of K_d for selenium equal to zero when the injected solution (with carriers) was used, and a value of 2.4 for selenium without carriers, cf Table IV.

Neodymium

Neodymium could not be detected during an observation time of 60 days. This is explained by the short half-life of ^{147}Nd and the high K_d values (Table IV), which means a strong retention of neodymium. Very high values for the neodymium K_d , in the range of 1 600 - 4 000 ml/g were obtained for granitic material (Allard, 1977). Our laboratory experiments showed that some neodymium precipitated, probably as hydroxide or carbonate, which will initially lower the concentration available for migration. A much larger activity has thus to be injected in order to demonstrate a complete concentration curve for neodymium in the fractures investigated.

Cesium

Cesium could not be detected during an observation time of 60 days. This is explained, in a similar way as for neodymium, by the high K_d value (Table IV) which means a retention for cesium sufficiently strong to allow the complete decay of the ^{131}Cs radioactivity before arriving at the observation boreholes.

4.7 Dispersion coefficient and mean transit time

The dispersion coefficient and the mean transit time for ^{82}Br and ^{85}Sr in the flow path B8 - B2 were calculated using the method given by Lenda and Zuber (1970). The calculations were based on the following assumptions: radial flow, instantaneous injection and infinite medium.

In Table V are listed calculated values of the transit time t_o , the dispersion coefficient D_x and the dimensionless parameter $D_x/v \cdot x$; v and x being defined in equation (3) below.

The dispersion coefficient and the mean transit time were also estimated for ^{82}Br and $^{99\text{m}}\text{Tc}$ in the flow path B8 - B7 using the same methods and assumptions as for the B8 - B2 flow path above. The values $3.6 \text{ m}^2/\text{h}$ and 18.8 h were obtained for both ^{82}Br and $^{99\text{m}}\text{Tc}$.

By solving the dispersion equation for the case of instantaneous injection, using the Laplace Transform method, the solution reads (Lenda and Zuber, 1970):

$$\frac{c'_x \cdot x}{m} = \frac{1}{\sqrt{4 \frac{D_x}{v \cdot x} \left(\frac{t}{t_o}\right)^3}} \cdot \exp \left[- \frac{(1 - t/t_o)^2}{4 \frac{D_x}{v \cdot x} \cdot \frac{t}{t_o}} \right] \quad (3)$$

where	c'_x	linear tracer concentration
	m	total mass (activity) injected
	D_x	longitudinal dispersion coefficient
	v	mean pore velocity of conveying fluid
	x	injection to observation distance
	t_o	mean transit time of tracer
	t	time variable

By inserting the calculated values for t_o and D_x into equation (3) the theoretical breakthrough curves were calculated and are compared to the experimental values in the figures 12 and 13. Fairly good agreement for the leading as well as for the tailing part was obtained for the ^{85}Sr curve, whereas for ^{82}Br there is good agreement only for the leading part of the curve.

5. BENTONITE EXPERIMENTS

The aim of this second experimental series was to test how effectively the injection of bentonite into rock fractures could further retard any released nuclides. To get comparable data bentonite was injected into the same fracture zones of borehole B8 (figures 2 and 3) as were utilized for the migration study. For the same reason previously studied radionuclides were used.

5.1 Bentonite injection procedure

Slurries of bentonite were injected between rubber packers into narrow sections (2 to 4 metres) of the borehole and with successively increasing concentration of bentonite and applied pressure (Nilsson, 1977). Totally about 2 000 kg (dry weight) of bentonite was injected. Of this amount, 1 400 kg, corresponding to about 20 m³ of bentonite slurries, was injected into the high permeability zone below 71 metres. In this case bentonite reached as far as to borehole B8 as well as to borehole B7.

In a final operation, bentonite was injected in high concentrations and with high pressure from the ground surface; i. e. the whole length of the borehole was included. Before the subsequent tracer tests the wall of the borehole was cleaned with a wire brush and rinsed by water flushing.

5.2 Radionuclide injection procedure

The radionuclides were injected into the lower part of the borehole, endeavouring to get them concentrated in the fracture zone at the 72 metres level. Before the injection of tracers a sealed plastic tube was placed in the borehole, in which a spectrometer borehole probe could be operated for control of any activity loss. The borehole loggings did thus not influence the water movements.

Two isotope tracers, ^{82}Br and ^{85}Sr , were used and injected in two separate runs. 3.2 mCi of ^{82}Br was first injected in order to check the injection procedure. After the decay of most of the ^{82}Br activity, 0.87 mCi of ^{85}Sr was injected. As all bromine activity was observed to remain in the borehole in the first run, indicating watertight fractures, the simultaneous injection of ^{82}Br (as a reference tracer for the water movement) was considered unnecessary in the second run. "During the summer 1978 measurements of the watertable were performed. The watertable was approx. +1.1 m in B2, + 1.4 in B7 and + 10.4 m in B8 (compare fig 2)."

Other experimental conditions were the same as in the previous migration study, e. g. water was pumped from borehole B2 at a constant rate of 0.1 litre/second.

5.3 Radioactivity measurements

The possible migration of tracer activity from borehole B8 was controlled in two ways:

- Samples of pumped water were regularly taken from the observation borehole B2 and analysed in the laboratory by similar methods to those used in the migration study.
- The remaining tracer activity in the injection hole B8 was selectively measured with a borehole gamma ray spectrometer; the ^{82}Br analysis based on the gamma energies 554 keV and the ^{85}Sr analysis on the 514 keV gamma-ray. Background activity, measured before the radionuclide injection, was subtracted from the bromine and strontium energy interval values.

The borehole spectrometer was energy-calibrated using suitable radioisotope sources. The calibration was checked before and after each logging run by recording characteristic energy peaks of the tracers. This is illustrated in fig 14, which shows the ^{85}Sr spectrum measured at the fracture zone level. The lower energy peak might be due to ^{75}Se which precipitated in connection with the previous migration study.

5.4 Results and discussion

Neither ^{82}Br nor ^{85}Sr could be detected in water samples from borehole B2. The detection limits for ^{82}Br and ^{85}Sr were 50 pCi/litre and 12 pCi/litre, respectively. These results were confirmed by the results of borehole measurements in the injection hole B8; no loss of tracer activity from the borehole could be observed. In Table VI are listed the results of successive borehole measurements of the integrated ^{82}Br activity in the depth interval 69 - 73.5 metres, which includes the highly permeable zone utilized for injection in the previous migration experiments but now filled up with bentonite.

The corresponding values of integrated ^{85}Sr activity for the interval 69 - 73.5 metres are listed in Table VII. In this case the observation time is considerably longer and therefore more interesting. Similarly to ^{82}Br , the variation of the values is within the estimated errors; typical sources of error being temperature effects on the borehole probe, background variation, statistical fluctuations of radiation, probe position determination, etc. Moreover, the strontium concentration in the borehole fluid could also change with time without any real loss of activity from the borehole, e. g. by diffusion and by sorption on bentonite, which is present in fissure openings, on the borehole wall and sedimented to the bottom of the hole. The concentration - depth curves of ^{85}Sr from two measurements with six months' time difference are shown in fig 15.

The migration rate observed for strontium after the bentonite injection is at least 135 times slower than the rate of the groundwater.

6. SUMMARY AND CONCLUSIONS

Nuclide migration in fractured rocks was studied in field experiments in a granite-gneissic outcrop within the Studsvik research centre area. Radionuclides representing long-lived fission products of the elements selenium, technetium, tin, cesium, iodine, neodymium and strontium were injected into a fracture zone intersecting one of the testholes at a depth of 72 metres below the ground surface. Concentration-time curves of activities arriving at the pumping borehole were measured. The distance between the boreholes was 51 metres and the pumping rate 0.1 litre/s. The mean transit time and the dispersion coefficient for water were calculated to be 57 h and $8.8 \text{ m}^2/\text{h}$, respectively. A shorter flow path (22 metres) could be utilized by intermittent sampling in a third borehole.

The radioactivity measurement techniques included gamma-spectrometric borehole logging, continuous activity measurements of pumped water and water sampling followed by laboratory analysis. Different combinations of radionuclides were injected simultaneously and selectively registered by gamma-ray spectrometry using a Ge(Li) detector in the field as well as in the laboratory.

Technetium and iodine travelled as anions with the same velocity as that of water. Strontium was retarded by a factor of about 6. Neither cesium nor neodymium could be detected. This is in agreement with their high retardation factors, estimated from mass distribution constants, determined in the laboratory. Selenium was partly and tin probably almost completely precipitated when injected into the borehole, which complicated the interpretation. A selenium pulse of short duration arrived simultaneously with the bromine pulse. About 20 days later the beginning of a second breakthrough of selenium and a first breakthrough of tin was observed in the shorter flow path.

In a second series of experiments it was shown that the fracture system surrounding the injection borehole could be effectively grouted with bentonite. Subsequent tracer tests with strontium-85 showed that after more than twelve months all the injected activity still remained in the borehole. It was found that bentonite injection is a very effective way of reducing the radionuclide migration through strongly fractured rock.

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TABLE I Activities of some important fission products in highly radioactive waste after 100 and 1 000 years.
From Kjellbert, N (1977)

Nuclide	T _{1/2} years	Activity (Ci/ton uranium) after	
		100 years	1 000 years
⁷⁹ Se	65 000	0.39	0.39
⁹⁰ Sr	28.1	6 500	-
⁹³ Zr- ^{93m} Nb	1.5 · 10 ⁶	1.9	1.9
⁹⁹ Tc	210 000	14	14
¹⁰⁷ Pd	7 · 10 ⁶	0.12	0.12
¹²⁶ Sn- ^{126m} Sb- ¹²⁶ Sb	100 000	0.57	0.56
¹²⁹ I	17 · 10 ⁶	0.038	0.038
¹³⁵ Cs	3 · 10 ⁶	0.25	0.25
¹³⁷ Cs- ^{137m} Ba	30	11 000	-
¹⁵¹ Sm	87	570	0.44

TABLE II Isotope tracer used

Element studied	Isotope used	Half-life	Gamma energies used in measurements	MPC - Public water
			keV	$\mu\text{Ci/cc}$
Sodium	^{24}Na	15 h	1 369	$3 \cdot 10^{-5}$
Selenium	^{75}Se	120 d	265	$3 \cdot 10^{-4}$
Bromium	^{82}Br	35 h	554 - 777	$4 \cdot 10^{-5}$
Strontium	^{85}Sr	64 d	514	$1 \cdot 10^{-4}$
Technetium	$^{99\text{m}}\text{Tc}$	6 h	142	$3 \cdot 10^{-3}$
Tin	^{113}Sn	115 d	392	$8 \cdot 10^{-5}$
Iodine	^{131}I	8 d	365	$3 \cdot 10^{-7}$
Cesium	^{131}Cs	9.7 d		$9 \cdot 10^{-4}$
Neodymium	^{147}Nd	11.1 d	531	$6 \cdot 10^{-5}$

TABLE III Combination of injected isotopes

Injection number	Isotopes injected	Chemical form (and amount of carriers)	Activity injected	Amount of solution
			mCi	litres
4	^{82}Br	NH_4Br (1 g)	2.44	
1977-02-09	$^{99\text{m}}\text{Tc}$	TcO_4^-	12.9	11.8
	^{131}I	NaI	0.92	
5	^{82}Br	NH_4Br (1 g)	2.4	
1977-03-11	^{75}Se	H_2SeO_3 (345 mg)	1.85	
	^{85}Sr	SrCl_2 (100 mg)	0.65	5.6
	^{113}Sn	SnCl_4 (158 mg)	0.90	
	^{147}Nd	$\text{Nd}(\text{NO}_3)_3$ (1 g)	2.15	
6	^{82}Br	NH_4Br (1 g)	1.66	
1977-03-22	^{24}Na	NaCl	1.65	5.6
	^{131}Cs	CsCl	5	

TABLE IV Laboratory determinations of the mass distribution coefficient, K_d

Element	Isotope used	K_d (ml/g)
Strontium	^{85}Sr	1.5 (5)
Neodymium	^{147}Nd	44 (4)
Iodine	^{131}I	0
Selenium	^{75}Se	2.4
Cesium	^{134}Cs	560 (2)

TABLE V Mean transit times (t_o), dispersion coefficient (D_x) and the parameter D_x/vx

Isotope	t_o hours	D_x m^2/h	$\frac{D_x}{v \cdot x}$
^{82}Br	57	8.8	0.152
^{85}Sr	398	16.2	0.28

TABLE VI Total activity of ^{82}Br in the depth interval 69 - 73.5 metres of borehole B8

Hours after injection	27	45	95	142	167
Activity, mCi	0.463	0.501	0.429	0.482	0.486

TABLE VII Total activity of ^{85}Sr in the depth interval 69 - 73.5 metres of borehole B8

Days after injection	43	47	57	62	68	111	229	374
Activity, mCi	0.336	0.361	0.379	0.345	0.362	0.362	0.366	0.389

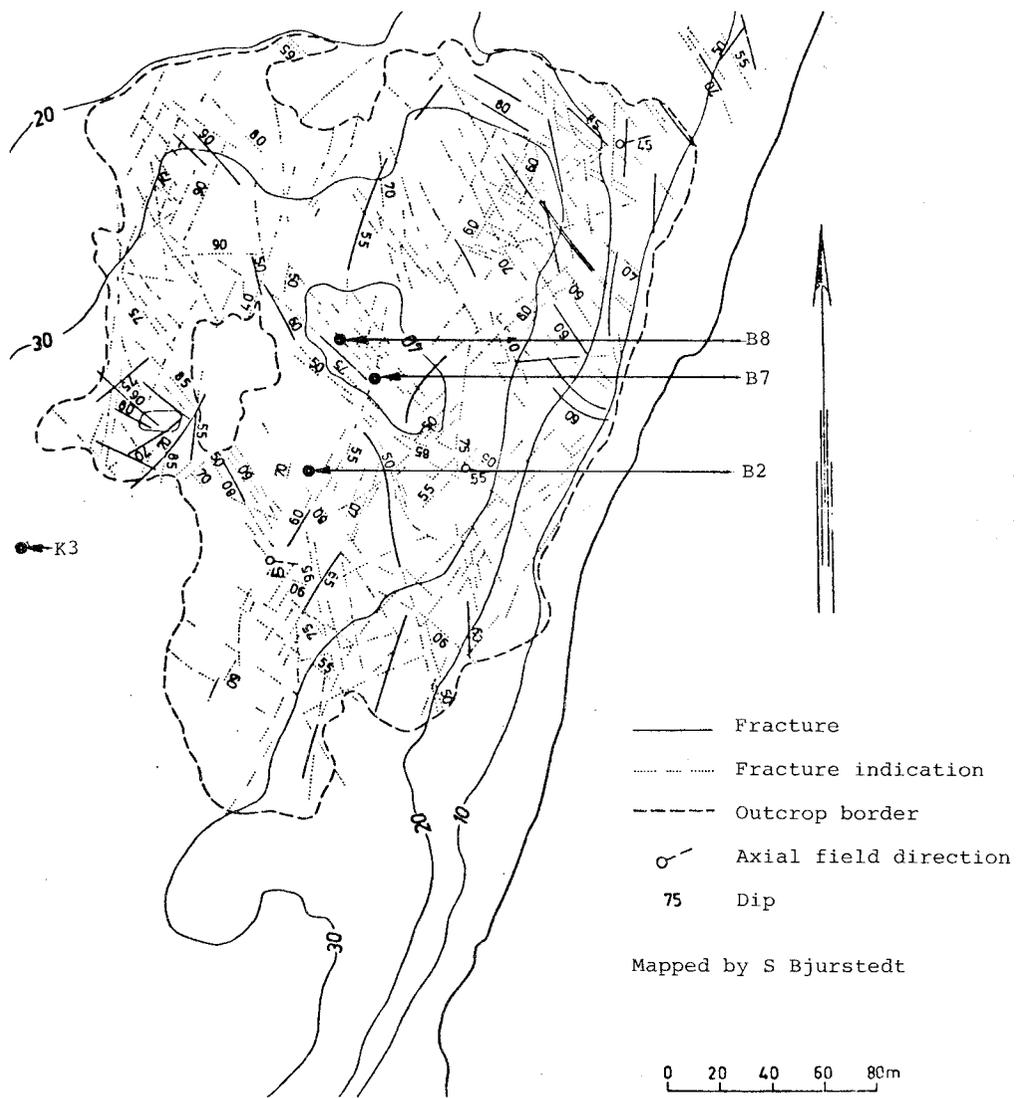


Fig 1. FRACTURE MAP OF THE EXPERIMENTAL AREA AND BOREHOLE POSITIONS

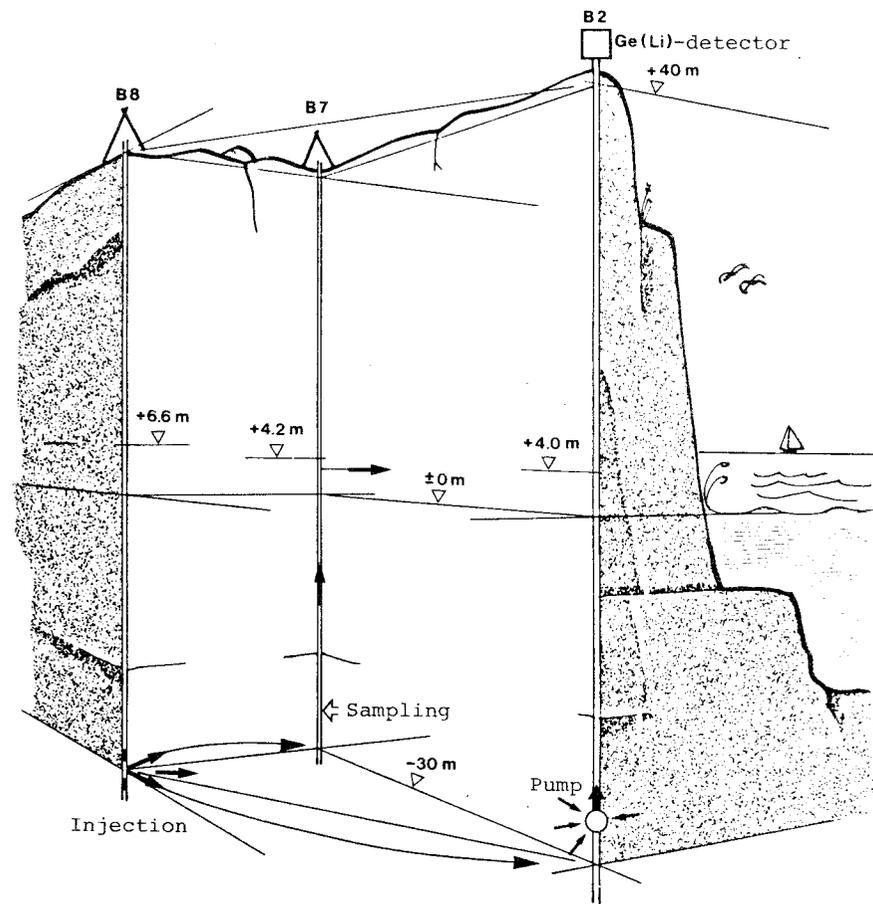


FIG 2. CUTAWAY VIEW OF THE "IN SITU" EXPERIMENTAL AREA

Sectional permeability

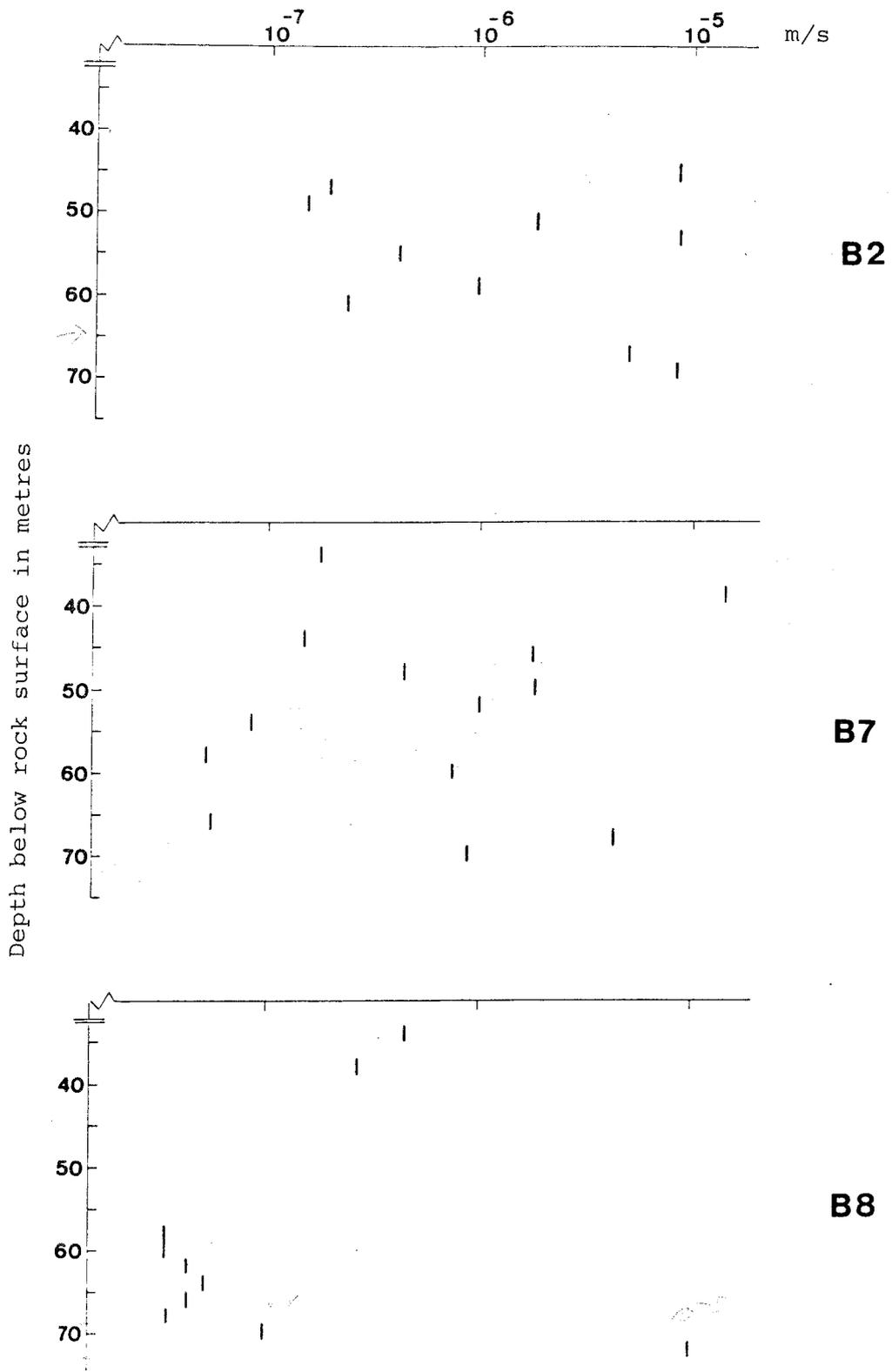


Fig 3 Permeability measurements of 2 metre sections in boreholes B2, B7 and B8. Limit of sensitivity: $2.5 \cdot 10^{-8}$ m/s.

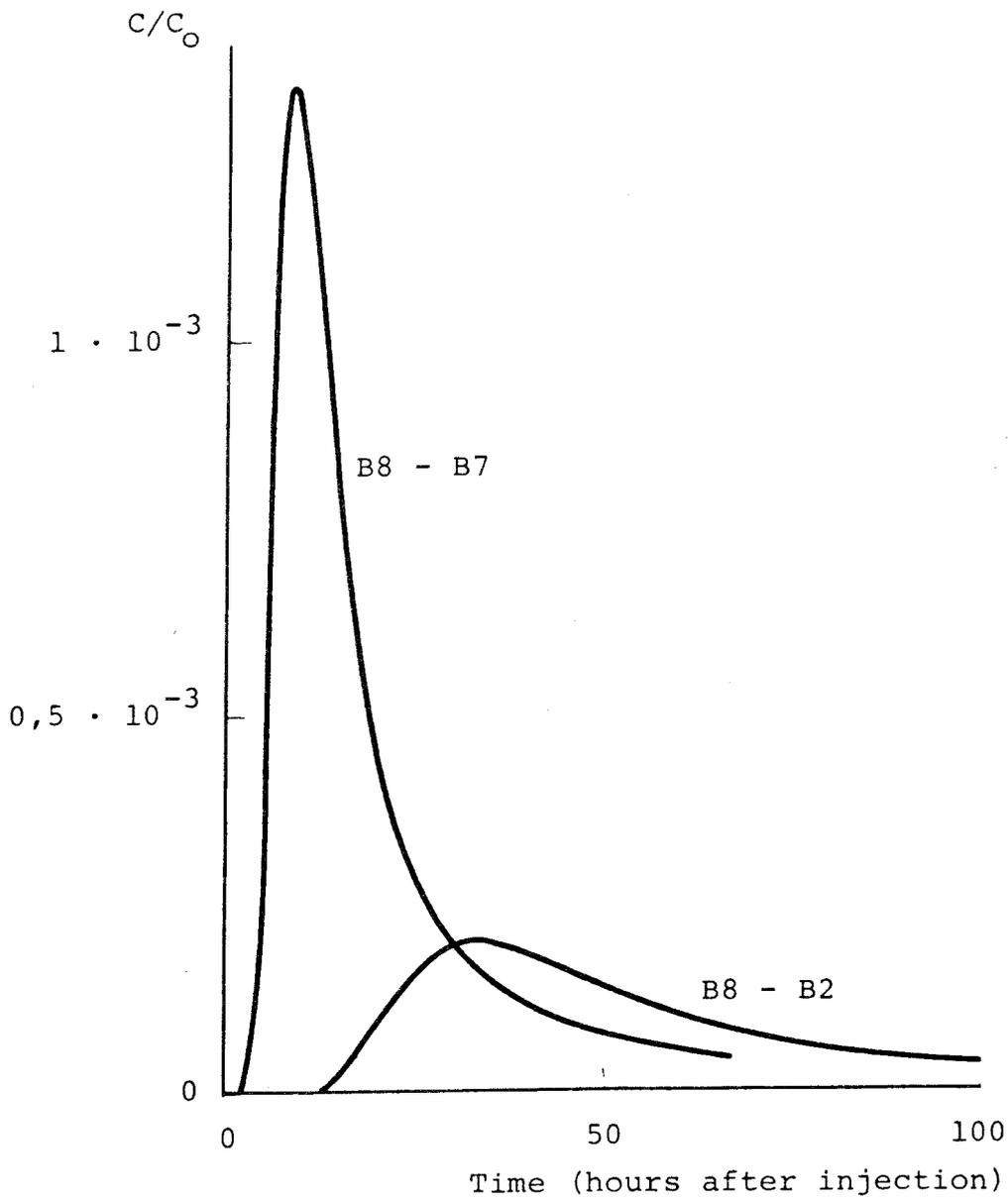


FIG 4. CONCENTRATION CURVES OF ^{82}Br FOR THE TWO FLOW PATHS

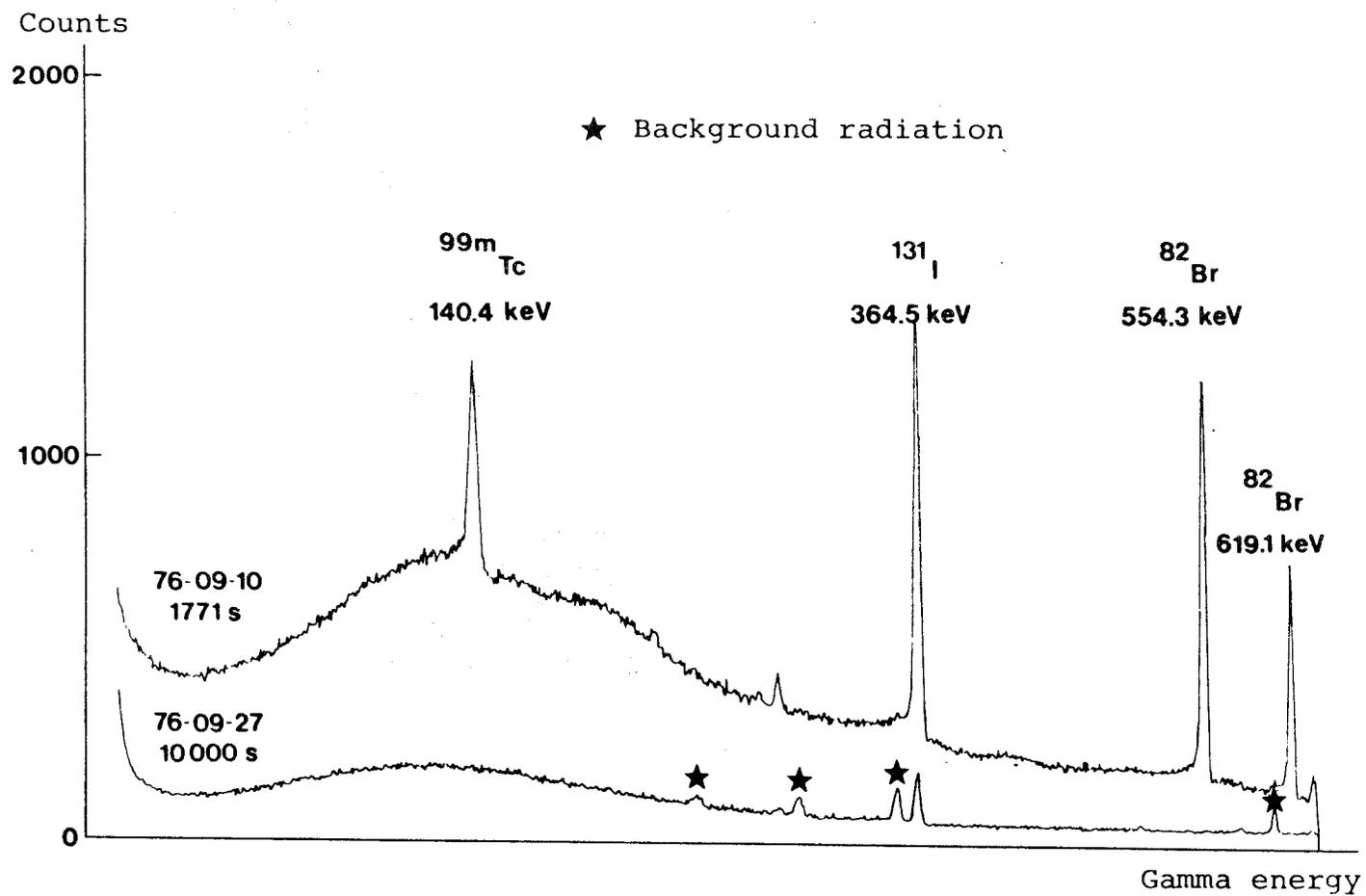


FIG 5. GAMMA SPECTRA, MEASURED WITH THE Ge(Li) DETECTOR AT BOREHOLE B2

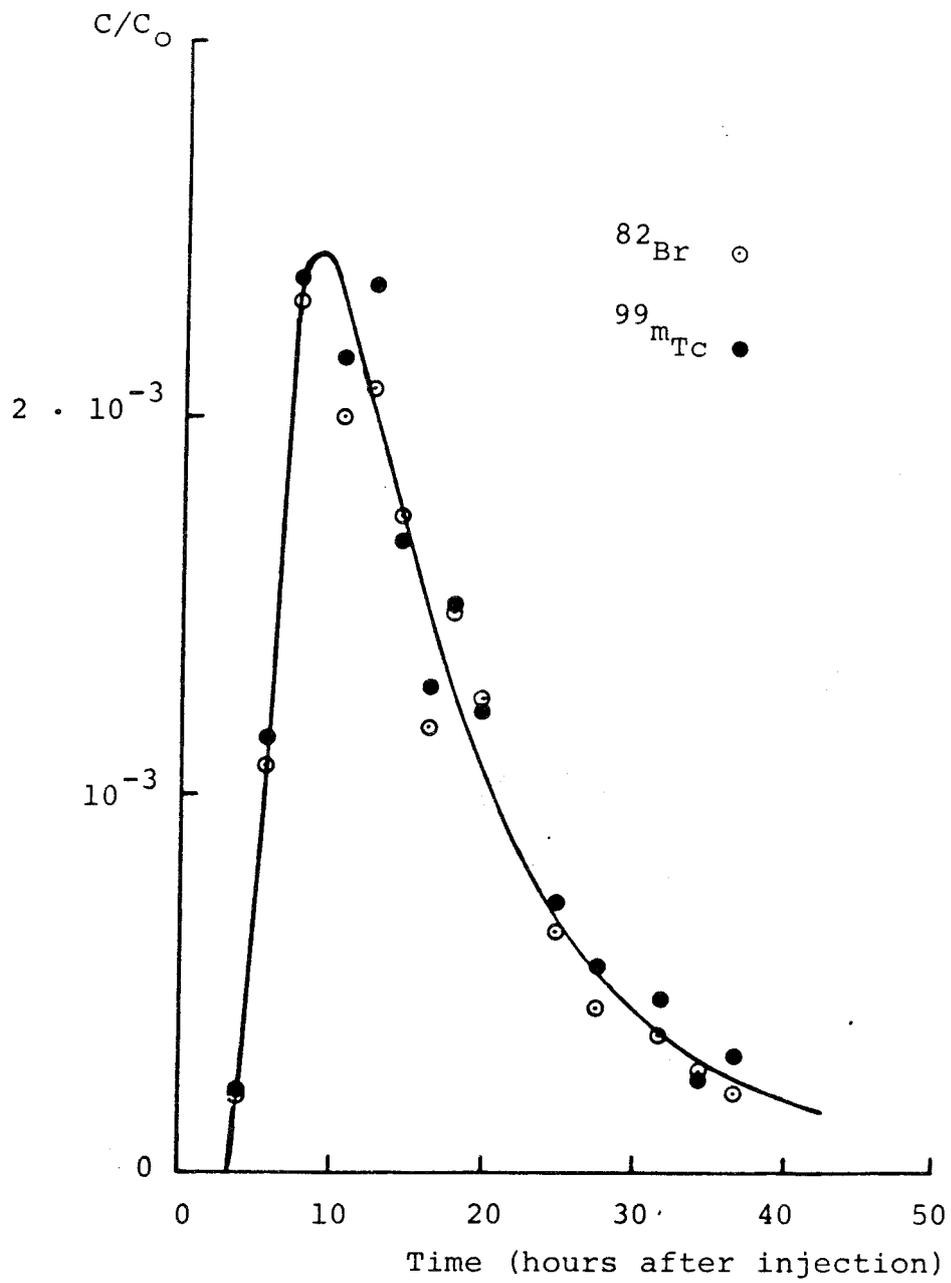


FIG 6. CONCENTRATION CURVES OF ^{82}Br AND $^{99\text{m}}\text{Tc}$
 FOR THE FLOW PATH B8 - B7

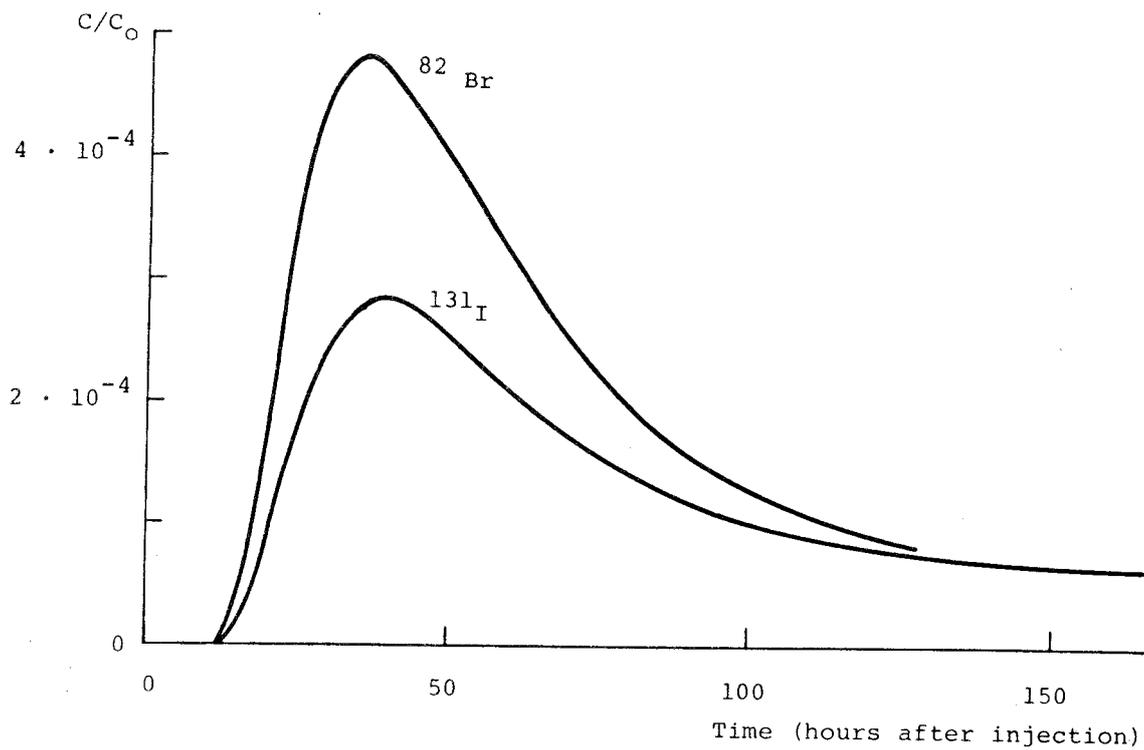


FIG 7. CONCENTRATION CURVES OF ^{82}Br AND ^{131}I FOR THE FLOW PATH B8 - B2

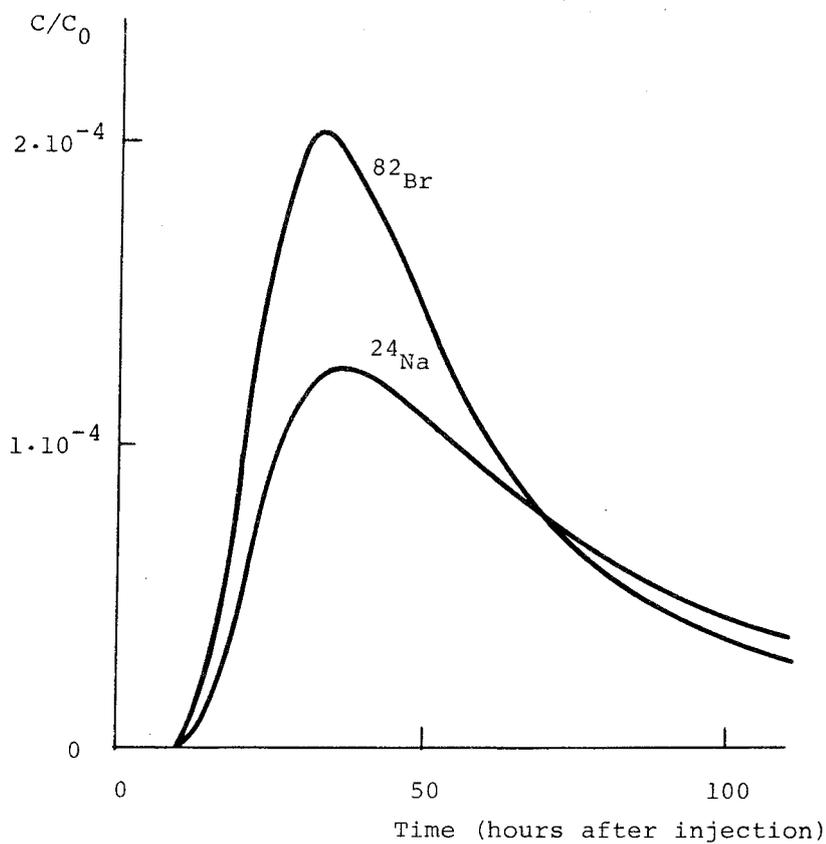


FIG 8. CONCENTRATION CURVES OF ^{82}Br AND ^{24}Na FOR THE FLOW PATH B8 - B2

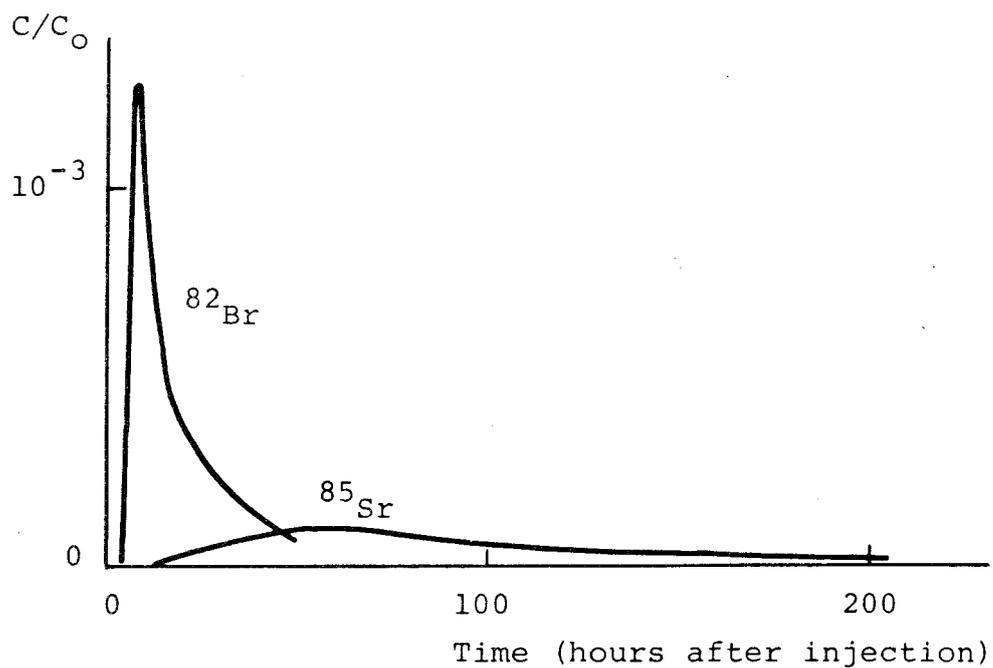


FIG 9. CONCENTRATION CURVES OF ^{82}Br AND ^{85}Sr FOR THE FLOW PATH B8 - B7

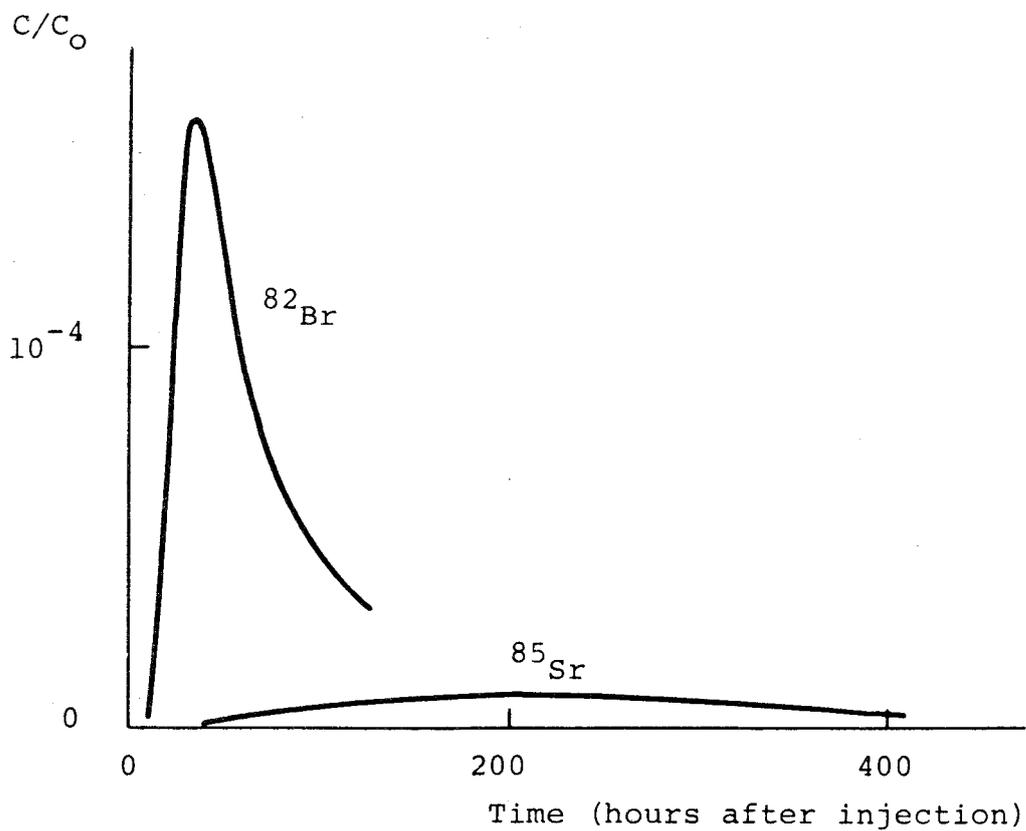


FIG 10. CONCENTRATION CURVES OF ^{82}Br AND ^{85}Sr FOR THE FLOW PATH B8 - B2

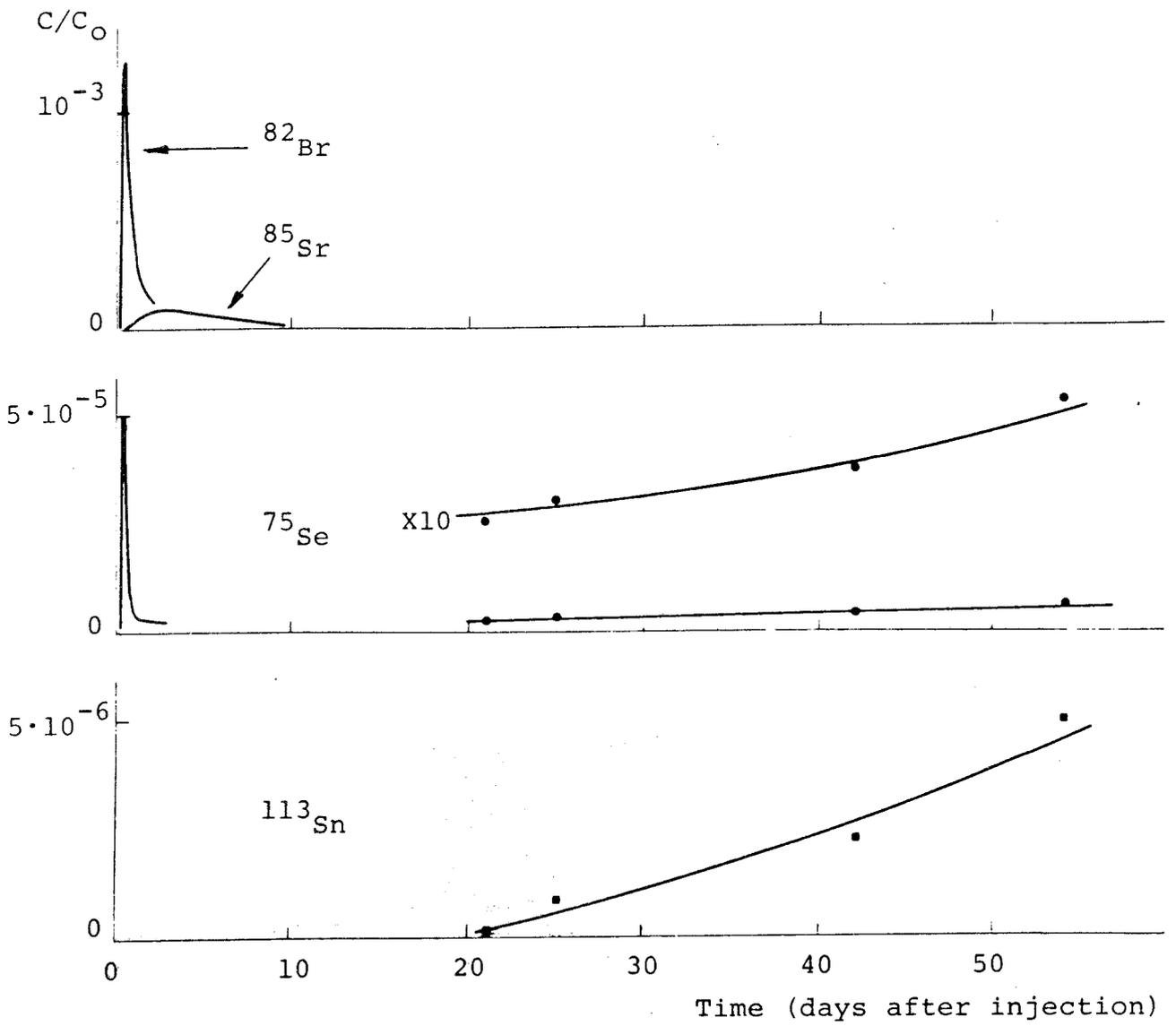


FIG 11. CONCENTRATION CURVES OF ^{82}Br , ^{85}Sr , ^{75}Se AND ^{113}Sn FOR THE FLOW PATH B8 - B7

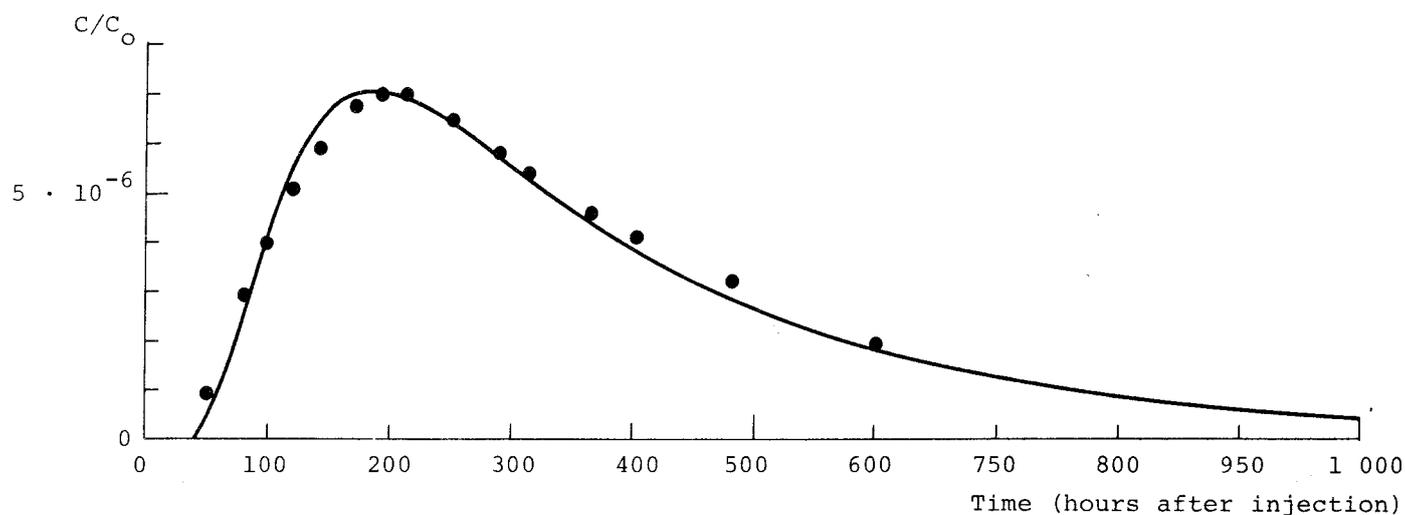


FIG 12. THEORETICAL CURVE AND EXPERIMENTAL POINTS FOR ^{85}Sr IN PATHWAY B8 - B2

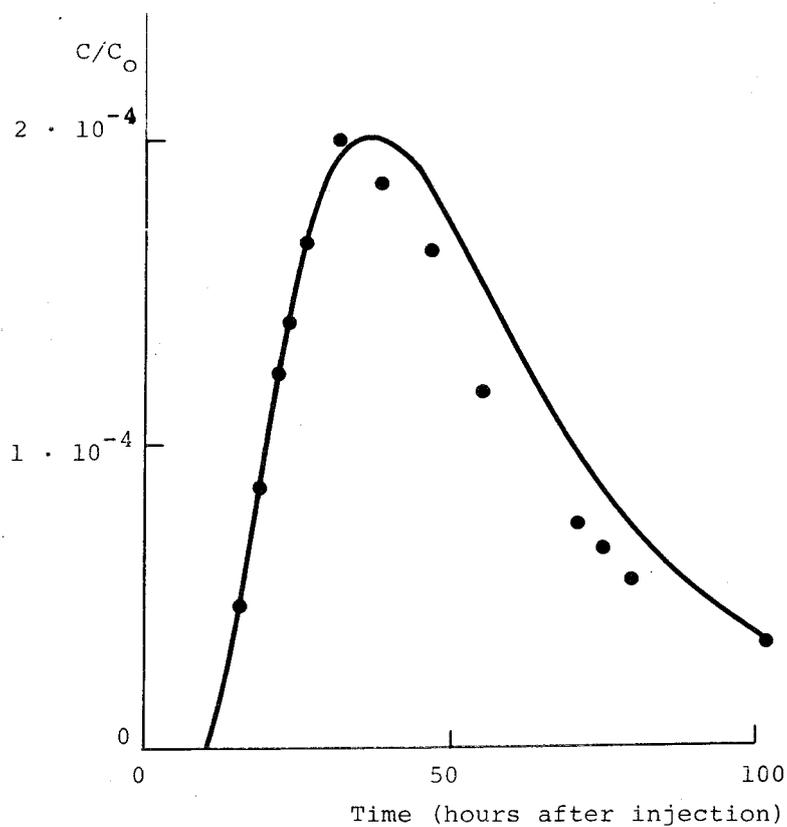


FIG 13. THEORETICAL CURVE AND EXPERIMENTAL POINTS FOR ^{82}Br IN PATHWAY B8 - B2

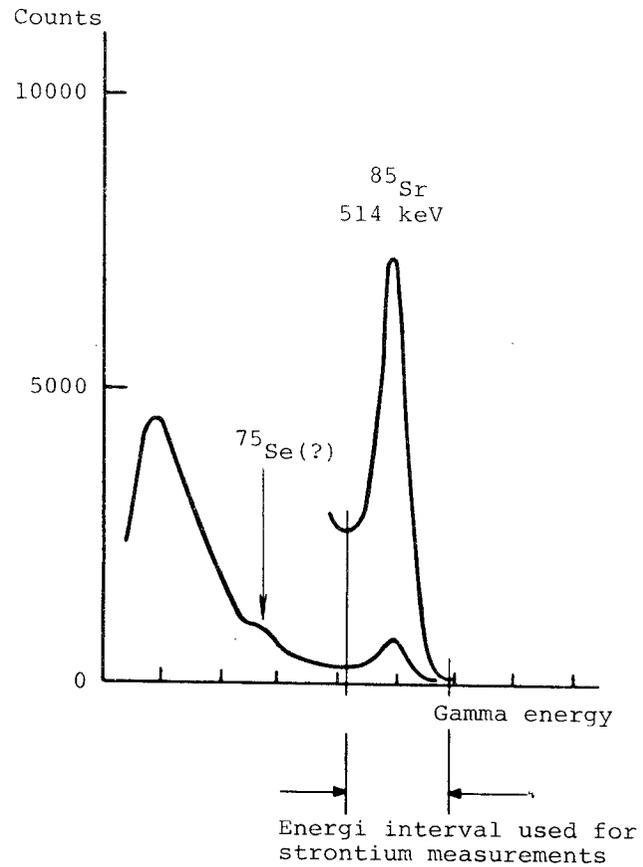


FIG 14. GAMMA SPECTRUM MEASURED AT 72 M LEVEL IN BOREHOLE B8. NaI DETECTOR Ø 1/2" x 1".

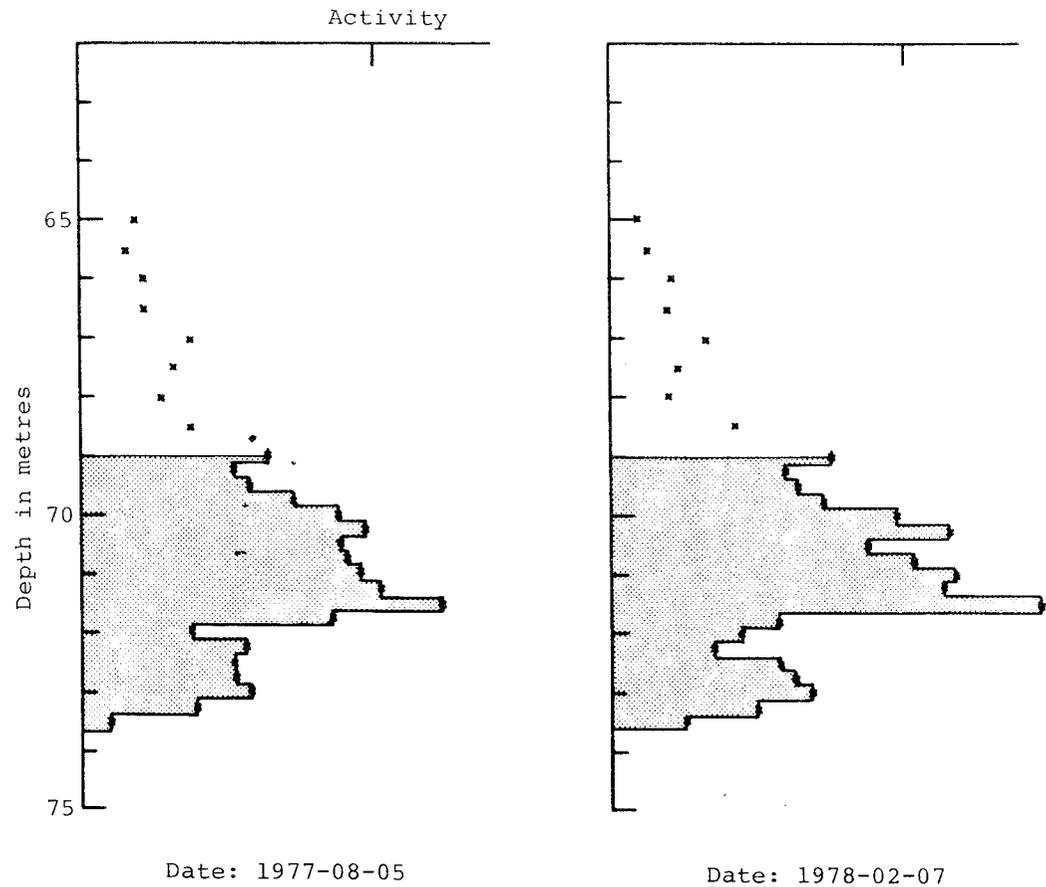


FIG 15. LOGGING DIAGRAM OF ^{85}Sr ACTIVITY (CORRECTED FOR RADIOACTIVE DECAY) IN BOREHOLE B8 AT TWO MEASUREMENT TIMES. SHADED AREAS: TOTAL ACTIVITY LISTED IN TABLE VII.

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