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June 2005

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This report concerns a study which was conducted for SKB. The conclusions and viewpoints presented in the report are those of the authors and do not necessarily coincide with those of the client.

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Abstract

This document is an account of the Multi-offset Multi-azimuth Vertical Seismic Profiling (VSP) survey performed in 2004 at Forsmark, in boreholes KFM01A and KFM02A, as well as of the processing, interpretation and results obtained.

Methods developed specifically for fracture detection and mapping in hard rocks were used for the processing and interpretation. The key elements of the multi-azimuth VSP data processing scheme are various Image Point (IP) filtering techniques, aimed at enhancing the reflected wavefields and at separating events generated by reflectors with different orientations. Polarization analysis and filtering were also used to enhance the P-wave reflected field and to suppress S-wave conversions. With the reflection events emphasized by IP filtering, 3-D orientations of the reflectors were determined interactively, with the help of clustering analysis.

Vibrometric carried out VSP surveys with time-distributed swept-impact sources instead of explosives since 2000. The VIBSIST-1000 source uses a tractor/excavator-mounted hydraulic rock-breaker, delivering approximately 2,000 J/impact at a mean rate of 10 impacts/second. The seismic signals are produced as rapid series of impacts, the impact intervals being monotonically increased to achieve a non-repeatable sequence. As the energy is built up from relatively low-power impacts, the high frequency components of the seismic signal are maintained. 3-component downhole geophones were used.

The objective of the VSP measurements has been to provide a better estimate of the dip of known fracture zones, in the area around boreholes KFM01A and KFM02A. In particular, the VSP measurements were meant to complete and confirm the orientation and position of the gently dipping site-scale reflectors observed in the previously performed surface seismic surveys and to identify and map possible steeply inclined features.

The surveys were conducted to a borehole length of 775 m in each borehole, the distances from the top of each borehole to the furthest shot points being around 1,200 m. Major seismic features have been mapped to depths in excess of 1.5 km and to a lateral distance of 2–3 km from the borehole collars. The frequency band was 30–250 Hz. The velocities of the P-waves were within $\frac{+}{-100}$ m/s of the 5,750 m/s average.

Reflectors were classified in three categories: strong, probable and possible, according to an evaluation scheme based on visibility, extent and consistency amongst several profiles. Most of the gently dipping reflectors were interpreted as dipping towards S-SE. Steeply dipping reflectors were assigned mainly to two groups, dipping towards SW and towards S-SE.

Sammanfattning

Denna rapport beskriver vertikal seismisk profilering (VSP) utförd i Forsmark 2004 i borrhålen KFM01A och KFM02A. Rapporten beskriver även processering och tolkning av data samt diskuterar resultaten.

I processeringen och tolkningen användes metoder specifikt utvecklade för att detektera och orientera sprickzoner i kristallint berg. De viktigaste stegen i sk multi-azimuth VSP är en spegelpunkt ("Image Point", IP) filtreringsteknik som syftar till att förstärka reflekterade vågfronter och separera signalen från reflektorer med olika orienteringar. För att förstärka de reflekterade P-vågarna och undertrycka S-vågs konversion användes polarisationsanalys och filtrering. Reflektorernas orientering i rummet bestämdes därefter interaktivt med hjälp av såkallad "clustering analysis".

Vibrometric har utförd VSP undersökningar med hydraulisk hammare sedan år 2000. Denna VIBSIST teknik består av en traktorgrävare med en hydraulisk hammare som presterar ca 2 000 J per slag med i genomsnitt ca 10 slag per sekund. Den seismiska signalen skapas genom en snabb serie av slag, där slagintervallet ändras monotont för att åstadkomma en lämplig sekvens. Eftersom den seismiska signalen byggs upp av slag med förhållandevis låg energi, bibehålls de höga frekvenserna. I borrhålen användes 3-komponent geofoner.

Syftet med VSP undersökningarna var att bestämma orienteringen av sprickzoner i området runt kärnborrhålen KFM01A och KFM02A. Speciellt viktigt var att säkerställa och bekräfta orienteringen och läget av de svagt stupande reflektorerna som observerades i tidigare utförda markbaserade reflektionsseismiska undersökningar, samt möjligheten att identifiera eventuella brant stupande reflektorer.

Undersökningarna utfördes ned till en borrhålslängd på 775 m i varje borrhål. Det största avståndet från borrhålens topp till skottpunkterna var 1 200 m. Signifikanta seismiska reflektorer har observerats ned till djup som överstiger 1 500 m och horisontellt upp till 2–3 km från borrhålen. Frekvensbandet var 30–250 Hz, och de observerade P-vågshastigheterna var inom intervallet 5 750 m/s \pm 100 m/s.

Reflektorerna klassificerades i tre kategorier: starka, sannolika och möjliga. Denna kategorisering baserar sig på visibilitet, utsträckning och konsistens mellan observationer från flera profiler. De flesta svagt stupande reflektorer tolkades som stupande mot S-SO. De brant stupande reflektorerna förekommer i två huvudgrupper, stupande mot SV respektive mot S-SO.

Contents

1 Introduction

This document reports the Multi-offset Multi-azimuth Vertical Seismic Profiling (VSP) survey performed in boreholes KFM01A and KFM02A in 2004, as well as the processing, interpretation and results obtained. The work was carried out following the Activity Plan AP PF 400-04-60. The documents describing the activity are listed in Table 1-1.

The locations of the boreholes and of the shotpoints are shown on the map in Figure 1-1. Figure 1-2 and Figure 1-3 display construction details of the boreholes.

Figure 1-1. Map of the Forsmark investigation area showing the location of the boreholes KFM01A (DS1) and KFM02A (DS2). The shot points are indicated with black dots.

Figure 1-2. Technical data for borehole KFM01A.

Figure 1-3. Technical data for borehole KFM02A.

Multi-offset VSP has been used, over the past 20 years, for mapping fractures and fracture zones with various hard rock applications, from nuclear waste disposal /Cosma et al. 2001a/, to ore delineation /Keskinen et al. 1999/, to rock engineering /Keskinen et al. 2000/. Fracture zones in hard rock may display a weak seismic contrast and specialist-imaging techniques must often be used to identify them. One method that produced good results in the past is based on computing time-depth functions for all possible reflectors in a given rock volume and investigating the coherency along each such path. Actual reflection events are emphasized in the process even if they do not display a remarkable amplitude standout in the raw profiles /Cosma, 1995, 1990/.

The VSP method provides a favorable geometry for mapping both steeply and gently dipping features /Cosma et al. 2001b/. Receivers located in the bedrock minimize the loss of resolution due to near-surface signal absorption.

Recent surveys /Juhlin et al. 2002, Cosma et al. 2002a/ have been carried out with a time-distributed swept-impact source, the VIBSIST /Enescu and Cosma, 1999; Cosma and Enescu, 2001/, instead of explosives. With this source, the seismic signals are produced as rapid series of impacts, the impact intervals being monotonically increased to achieve a non-repeatable sequence. As the energy is built up from a large number of relatively lowpower impacts, the high frequency components of the seismic signal are maintained.

Methods developed over the past 10 years /Cosma et al. 1994a, 1994b; Heikkinen et al. 1994, 1995/ and recently summarized in /Cosma and Enescu, 2002b/ were used for the processing and interpretation, e.g. the interactive determination of the orientation of the fracture zones (reflecting interfaces) using simultaneously several VSP boreholes.

The data processing usually starts with procedures aimed at reducing or eliminating tubewaves and direct P and S onsets. Median and band-pass filters are used for this purpose. Further processing consists mainly of Image Point (IP) filtering techniques, aimed at enhancing the reflected wavefields and at separating events generated by reflectors with different orientations. With the IP transform, introduced by /Cosma, 1990; Cosma and Heikkinen 1996/, stacking is performed along hyperbolic paths corresponding to the timedepth functions of possible reflectors. Due to this "natural" choice of the stacking paths, the coherency can be used effectively to enhance the weak reflections. When all the profiles have been processed and the reflection events emphasized by IP filtering, 3-D orientations of the reflectors are determined interactively with the help of statistical methods, e.g. clustering analysis.

2 Objective and scope

The general goal of the VSP measurements has been to provide a better estimate of the dip of known fracture zones, identified by other methods, in the area around boreholes KFM01A and KFM02A. The expected result is a better understanding of the geometry of the zones.

In particular, the VSP measurements were meant to complete and confirm the orientation and position of the gently dipping site-scale reflectors observed in the previously performed surface seismic surveys and to identify and map possible steeply inclined features.

The surveys were conducted in boreholes KFM01A and KFM02A, to a maximum depth of 775 m in each borehole. As a rule of the thumb, sub-horizontal and gently dipping seismic targets can be mapped to a depth of one and a half times the depth of the boreholes and steeply dipping targets to a lateral distance of the order of the depth of the boreholes. However, this rather conservative estimate can roughly be increased twofold if the same investigation volume can be imaged from several boreholes and the seismic targets are nearly planar and continuous over distances comparable with the borehole depth. Hence, in the case of KFM01A and KFM02A, the imaging of nearly planar and extensive seismic features has been attempted to depths in excess of 1.5 km.

To aid the estimation of the azimuth of the seismic imaging targets, the measuring routine attempts to achieve a consistent directional response of the 3-component receivers by keeping them clamped to the borehole in the same positions while the source moves from one shotpoint to another. However, as two sources were used, the shot points have been assigned to two groups, each group containing at least four shotpoints. The receivers were unclamped and moved only when all shotpoints in both groups had been recorded.

3 Equipment

3.1 The seismic source – VIBSIST-1000

A time-distributed seismic source, the VIBSIST-1000, was used for the VSP surveys. The VIBSIST-1000 is based on the Swept Impact Seismic Technique (SIST), /Park et al. 1996; Cosma and Enescu, 2001/.

At each shot point, the VIBSIST source was activated for a period of 15–20 seconds during which time the impact frequency has been varied to generate a swept impact sequence. The sweep was repeated five times and, when needed, up to ten times to obtain a high signal-tonoise ratio.

The SIST concept requires a pilot signal to be measured by a sensor placed on the source or in its immediate vicinity, conveyed to the recording station and recorded together with the signals arriving from the receivers. Pilot signal cables were drawn from each shot point to the receiver borehole.

The VIBSIST-1000 source uses a tractor/excavator-mounted hydraulic rock-breaker, powered through a computer controlled flow regulator. The Forsmark VSP survey was performed with two sources, one mounted on a 12 ton Huddig 960 tractor and another on an 18 ton Atlas 1504 excavator. These were equipped with hammers delivering 1,500 and 2,500 J/impact at 400–800 impacts/minute. The computer-controlled flow regulator produced about 100 impacts per sweep.

The source coupling was generally similar from one shotpoint to another being however occasionally influenced by local variations of the ground conditions. A larger number of sweeps (up to 10) was used to compensate for poorer ground conditions. Interpretable data was acquired at each shot location, except shot point no 5 for KFM02A, which initially produced poor results. It was found that underground facilities exist in that region and the shot point was abandoned. To allow the mapping of structures dipping in various directions, the shot points were located at various azimuths around the VSP boreholes (see Figure 1-1, where the shots are marked with black dots and labeled in red).

Figure 3-1. The VIBSIST-1000 seismic sources used as a seismic source at Forsmark for the 2004 VSP surveys.

3.2 The seismic receivers – the Vibrometric R8-XYZ-C geophones

A 3-component geophone chain was used for the VSP measurements. The R8-XYZ-C geophone chain is equipped with 24 28-Hz geophones placed in eight 3-component modules (Figure 3-2). The z-component is directed along the hole and the x- and y-components are perpendicular to the z-component and to each other. The distance between the modules is five metres. The geophone chain contains also a down-hole preamplifier for each channel. The gain of the amplifiers is fixed at 34.2 dB. The frequency range is from 40 to 1,000 Hz. The overall sensitivity of the geophone-amplifier combination is 133 V/cm/sec. The units are equipped with side arms for clamping, activated by DC motors. The clamping control is independent for each unit.

Figure 3-2. Picture of the R8-XYZ-C geophone receiver chain and detail of a 3-component module with clamping arm.

3.3 The recording station

A PC-based acquisition system was used (Figure 3-3), consisting of an ICS 32-channels, 24-bits acquisition board, with LabView based acquisition interface.

The PC-based solution has been chosen instead of a standard 24-bit seismograph because of the unusually long records needed (100,000 samples/channel for a 20 s sweep at 0.2 ms sampling rate), which is unavailable with standard seismographs. The customwritten software allows concurrent signal processing, needed for on-line checking of the signal quality. The decoded seismic record is built by 32-bit stacking of approximately 1,000 individual pulses. The data is recorded on an internal hard disk in the ACH Vibrometric format.

Figure 3-3. The acquisition station used for the surveys at Forsmark.

4 Execution

4.1 General

The VSP investigations were carried out during 2004 in boreholes KFM01A and KFM02A. Figure 4-1 shows the layout of the boreholes and of the shot point locations. There are 20 shot point locations that were used around two boreholes. Technical details on these boreholes are given in Figure 1-2 and Figure 1-3.

Figure 4-1. Layout of boreholes and shot points used for the VSP investigations (see also Figure 1-1).

The starting directions and positions of the VSP boreholes, in site coordinates, are given in Table 4-1.

The VSP survey configurations are shown in Table 4-2. The coordinates of the shot points used for each VSP borehole are given in Table 4-3, and Table 4-4. The zero-offset (nearest to the borehole top) shot point for each VSP survey is printed bold.

The distances from the top of the boreholes to shot points used for the VSP surveys range from 38 to 1,168 m for KFM01A and from 23 to 1,234 m for KFM02A. The distances from shot points to deepest receivers in the boreholes range from 771 to 1,474 m.

Table 4-1. The maxibor coordinates of the boreholes KFM01A and KFM02A. The coordinates system is RT90-RHB70. The dip is measured from horizontal and the azimuth is measured clockwise from North.

Boreholes used for the 2004 VSP surveys	Coordinates for the top of the borehole			Starting direction		Coordinates at the bottom of the borehole		
	Northing X (m)	Easting Y (m)	Elevation Z (m)	Dip (°)	Azimuth (°)	Northing X (m)	Easting Y (m)	Elevation Z (m)
KFM01A	6699529.8	1631397.1	3.12		-84.7 318.35	6699628.8	1631258.4	-982.3
KFM02A	6698712.5	1633182.8	7.35		-85.4 275.70	6698764.9	1633088.9	-988.7

Shot point	Northing X (m)	Easting Y (m)	Elevation Z (m)
SP01	6698694.7	1633167.8	5.3
SP02	6698930.8	1633422.6	1.5
SP03	6699416.9	1633326.4	3.1
SP04	6699683.5	1632940.5	2.5
SP05	6699724.2	1632476.3	2.0
SP06	6698040.6	1632613.9	4.2
SP07	6698112.9	1632967.0	22
SP ₀₈	6698488.3	1633328.7	4.3
SP ₀₉	6698121.7	1633272.2	3.7
SP ₁₀	6698000.7	1633543.0	5.3

Table 4-4. The coordinates of the source positions for the VSP survey in KFM02A.

4.2 Preparations

The field campaign started on August 2nd, 2004 and carried on until August 22nd, 2004. An introductory meeting was held during the first day, followed by field reconnaissance and identification of the shot point locations. Trigger cables were spread first on the North side and then on the South side, for all shot points used with borehole KFM02A. Cable continuity was checked for all cables. The acquisition of VSP data proceeded first with only one Vibsist source since the second tractor was not available at the time. Full production with two tractors commenced several days later. Once the acquisition of all VSP profiles was completed for borehole KFM02A, the work proceeded for borehole KFM01A.

4.3 Execution of field work

The required number of sweeps used at each shot point was continuously re-evaluated as the survey progressed. The standard quality control procedure consisted of evaluating the signal-to-noise ratio of the first sweep and based on this observation a choice of 5 or 10 sweeps was determined. The quality control procedure also included the inspection of the last sweep recorded to ensure that the equipment had functioned properly. The number of sweeps was also increased as a function of the shot-receiver distance.

Following each data acquisition day, the raw data files were backed-up and deconvolved from long sweeps to primary (short) raw data files.

5 Processing and interpretation of the VSP data

5.1 VSP data processing

The processing sequence aims to improve the signal-to-noise ratio, so that the later events, e.g. reflections, become visible. As the reflection coefficients are generally low, the reflectors cannot normally be identified by amplitude standout alone. Phase consistency has also been found to be a sensitive indicator.

The processing sequence consists of:

- \rightarrow Frequency filtering;
- \rightarrow Picking of the first arrival times and velocity analysis;
- \rightarrow Rotation of the horizontal components to radial (R) and transverse (T), where "radial" stands for the direction perpendicular to the hole and pointing towards the source and "transverse" for the direction perpendicular to the radial and to the hole;
- \rightarrow Eliminating such wave-fields as the direct P, direct S and tube-waves, so that the reflections become visible. The direct wave fields and other coherent disturbances are removed by slant median filtering and the signal levels are adjusted in such a way that the amplitudes of different traces and different parts of the same trace become comparable;
- \rightarrow Image Space (IP) processing, which is done to enhance the reflected wave fields and separate reflection events originating at interfaces with different orientations;
- \rightarrow Image Space non-linear enhancement of reflected energy;

Once all the profiles have been processed and the reflection events emphasized, the positions of the reflectors are determined by interactive interpretation.

All data processing has been realized using Vibrometric proprietary software, which contains several unique routines developed, over the past two decades, specifically for multi-offset VSP imaging in hard rocks.

5.1.1 Data quality and frequency analysis

The data has been inspected for possible malfunctions of the measuring system, unusually high noise levels, possible errors in coordinates, time delays and trace order. At the end of this process all errors, if any, falling in the categories listed above were corrected.

The following operations were completed during this processing stage:

- The geometry information was incorporated into the data profiles.
- The data was band-pass filtered in the frequency band of the seismic signal. The frequency band selected for the filter is as wide as possible, the rejected spectral components corresponding to clearly identified sources of noise. The frequency band of the P-waves was estimated to be 30–250 Hz. A zero-phase band-pass filter [30–250] Hz was used to filter all data profiles. The frequency content of the recorded data is illustrated in Figure 5-1, which displays the spectra of the data profiles shown in Figure 5-4 (left) and Figure 5-8 (right).

Figure 5-1. Spectra of the axial, Z-component, profile measured from SP03 at KFM01A (left) and from SP02 at KFM02A (right).

• The signal levels were adjusted so that the average amplitudes of different traces and different parts of the same trace become comparable.

5.1.2 Rotation of horizontal components

The orientation of the horizontal components $(X \text{ and } Y)$ is not set or determined during the measurements and, due to the free rotation of the down-hole probes while changing position, the horizontal components show a poor trace-to-trace consistency in both amplitude and phase.

The rotation of the horizontal components is done computationally, assuming that the direct P wave is polarized along the source-receiver line. The X-Y trace pair is rotated so that after rotation the X component acquires the most possible of the P-wave energy and becomes the "Radial" component, while the Y component contains the minimum possible of the P-wave energy becomes the "Transverse" component. The Z-component remains directed along the borehole and it becomes the "Axial" component.

The radial, transverse and the axial components from KFM01A shot point SP03 are presented as an example in Figure 5-2, Figure 5-3 and Figure 5-4. KFM01A SP03 is a zerooffset shot point; see Figure 4-1. The direction of propagation from the source to receivers is along the borehole and the transverse component aligning procedure described above, based on the direct P-wave polarization cannot be applied.

As the receivers were not moved between the shots in the same group, the aligning has been done by extrapolating the polarization analysis made on the far-offset shots.

One does notice amplitude predominance of the direct P-wave arrivals in the axial (Z) component, Figure 5-4, while S-waves and PS-conversions predominate in the transverse components R and T. These conversions can provide certain structural information, as detailed in Figure 5-5.

Figure 5-2. Raw data profile KFM01A, shot point SP03, radial (R) component.

Figure 5-3. Raw data profile KFM01A, shot point SP03, transverse (T) component.

Figure 5-4. Raw data profile KFM01A, shot point SP03, axial (Z) component.

Figure 5-5. Raw data profile KFM01A, shot point SP03, transverse (T) component. The event marked with a cyan line corresponds to a P-S conversion (Vp = 5,750 m/s to Vs = 3,450 m/s) generated by a gently dipping interface (Dip = 20°), which intersects the borehole at 220–240 m.

The radial, transverse and axial components from KFM02A shot point SP02 are presented in Figure 5-6, Figure 5-7 and Figure 5-8. SP02 is a relatively near (but not zero) offset shot point. However, there is clearly observable direct P-wave energy in the radial component and corresponding direct S-wave energy in the transverse component, see Figure 5-6 and Figure 5-7, respectively. P-S wave conversions are also visible and can be associated with structure, as explained in Figure 5-9.

5.1.3 Velocity determinations

The velocity fields for all profiles measured in boreholes KFM01A and KFM02A show no abnormal characteristics. The interval velocities (slope of the trajectory length vs. travel time) remain within $\frac{+}{-100}$ m/s of the 5,750 m/s average for all shot points.

5.1.4 Suppression of direct P-wave and S-waves

An efficient way to eliminate direct P-wave and/or S-wave arrivals is by means of median filters or adaptive median filters taken along the slopes corresponding to $+/-$ the P-wave and/or S-wave velocity. The trajectory along which median filtering is done is not straight, but follows the true travel time curve. The direct P-wave energy along 5,750 m/s paths was removed from the VSP profiles, followed by the removal of the direct S-wave energy, along 3,450 m/s paths.

Figure 5-6. Raw data profile KFM02A, shot point SP02, radial (R) component.

Figure 5-7. Raw data profile KFM02A, shot point SP02, transverse (T) component.

Figure 5-8. Raw data profile KFM02A, shot point SP02, axial (Z) component.

Figure 5-9. Raw data profile KFM02A, shot point SP02, radial (R) component. The event marked with a cyan line corresponds to a P-S conversion (Vp = 5,750 m/s to Vs = 3,450 m/s) generated by a gently dipping interface (Dip = 20°), which intersects the borehole at 330–340 m. The event marked with a blue line corresponds to a P-S conversion (Vp = 5,750 m/s to Vs = 3,450 m/s) generated by a gently dipping interface (Dip = 20°), which intersects the borehole at 90–100 m.

5.1.5 Amplitude compensation and equalization

Amplitude compensation (AGC) is performed to cancel the effects of geometrical spreading and attenuation and to reconstruct the original amplitude variations along the trace. A variable gain operator is run along the records to increase the amplitude of later events assumed to have traveled along longer paths. The amplitude compensation for all three components was done with the same operator, estimated in a 50 ms window, so that the amplitude ratio between the components is conserved after AGC.

The pre-processed profiles from borehole KFM01A, shot point SP03 are presented in Figure 5-10, Figure 5-11 and Figure 5-12. Weak coherent reflection patterns start to be visible after pre-processing.

5.1.6 Image point filtering and reflector enhancement in the image space

One of the properties of the Image Point transform /Cosma, 1995; Cosma and Heikkinen, 1996/ is that, if the velocity field is correctly modeled, the coherent energy adds in phase (along hyperbolic paths), producing well-defined maxima in the IP (Image Point) space. This offers possibilities for advanced intricate processing; including enhanced polarization analysis, azimuth and dip filtering, as well as non-linear coherency-enhancement schemes.

A novelty of the processing of the Forsmark VSP data was that the Image Point processing was applied independently on panels of 27 traces instead of whole panels containing 136 traces. This was done in order to allow potentially discontinuous seismic targets to be shown as such, without forcing continuity by coherency enhancement over the whole depth range of the shot gather profiles.

Figure 5-10. Pre-processed profile KFM01A, shot point SP03, radial (R) component.

Figure 5-11. Pre-processed profile KFM01A, shot point SP03, transverse (T) component.

Figure 5-12. Pre-processed profile KFM01A, shot point SP03, axial (Z) component.

The data, from SP03, borehole KFM01A, shown in Figure 5-13, Figure 5-14 and Figure 5-15 present the profiles presented in Figure 5-10, Figure 5-11 and Figure 5-12 as enhanced by Image Point filtering using 27-traces panels.

The polarization information was also used to enhance the P-wave reflected field and to suppress S-wave conversions, which would be polarized perpendicularly to the direction of the P-waves. A brief description of the IP transform and the use of polarization analysis in the IP space are given in Appendix A. A non-linear coherency enhancement scheme (in the IP space) was also applied, the results for SP03, borehole KFM01A, being shown in Figure 5-16, Figure 5-17 and Figure 5-18.

Figure 5-13. IP-filtered profile from KFM01A, shot point SP03, radial (R) component.

Figure 5-14. IP-filtered profile from KFM01A, shot point SP03, transversal (T) component.

Figure 5-15. IP-filtered profile from KFM01A, shot point SP03, axial (Z) component.

Figure 5-16. Enhanced IP-filtered profile from KFM01A, SP03, radial (R) component.

Figure 5-17. Enhanced IP-filtered profile from KFM01A, SP03, transversal (T) comp.

Figure 5-18. Enhanced IP-filtered profile from KFM01A, SP03, axial (Z) component.

A few conclusions can be drawn by comparing the groups of graphs for raw data (Figure 5-2 to Figure 5-4), the preprocessed data (Figure 5-10 to Figure 5-12), the linear IP transform filtered data (Figure 5-13 to Figure 5-15) and the non-linearly filtered data (Figure 5-16 to Figure 5-18). These processing steps are subsequent, each being applied on the result of the previous. The pre-processing, which is a collection of largely standard operators, successfully eliminated the direct P- and S-waves but the low apparent velocities of some of the coherent patterns emerging after the processing sequence indicates their S-wave nature. Steeper reflectors may generate S-wave conversions with steeper apparent

velocities, these being potentially mistakenly interpreted as genuine P-reflectors, with spurious orientations. The IP transform polarization filtering eliminates most of the S-wave conversions. Non-linear IP filtering conserves the major identifiable coherent patterns, i.e. reflections, and makes them easier to identify in the subsequent interactive interpretation stage. Therefore, non-linear IP filtering is in fact a means of data reduction meant to alleviate interpretation.

6 Interpretation of the VSP data measured from boreholes KFM01A and KFM02A

Due to the wide diversity of reflection angles, to the local variations of reflectivity and, generally, to the relatively weak seismic response of reflecting interfaces such as faults and fracture zones, the "importance" of a reflector cannot reliably be quantified based solely on the amplitude of the corresponding event in a time-depth profile. The VSP profiles measured in crystalline rock tend to display a relatively numerous collection of reflection events with similar amplitudes.

The "importance" of a reflector is assessed by evaluating its consistency amongst several profiles. Consistency may be interpreted as an indication of spatial extent and continuity. Each depth-time profile is a 2-D representation and does not contain sufficient information for the complete determination of the 3-D orientation and position of a reflector. The interpretation process consists of finding the link between reflected events likely to represent the same feature in different profiles. Several profiles need to concur to resolve the orientation problem /Cosma and Enescu, 2002b/.

The reflector fitting procedure used with the Forsmark data is interactive and consists of superposing travel time functions on the reflection events. The travel time curves are displayed as coloured lines in Image Point filtered sections, see Figure 6-1. The depth of intersection in the receiver borehole (or its geometrical extension) and a dip – dip direction function best fitting the chosen trajectory are calculated. The dip and dip direction are varied according to the relation given by the fit and displayed as a stereographic projection. The dip and dip direction are determined by fitting amongst at least three different profiles, see Appendix A4.

For profiles measured in the same borehole the mean orientation of a reflector can be found in a stereographic representation (Wulff diagram – see Figure A-3 for a definition of the Wulff projection) at the intersection of the dip-dip direction curves computed for each profile, as shown in Figure 6-2. Each dip-azimuth curve (red) describes the possible orientations associated with the same time-depth function in a given profile. As fracture zones develop more or less along planes, their images in profiles measured with the same receiver layout are assumed to have the same intersection position along the layout and the stereographic projection is computed at this specific position. The intersection, the dip and the dip direction determine completely the position and orientation of a reflector. As the intersection is not exactly one point but a region, the "centre" of this region is determined by clustering analysis. The travel time curve computations are then performed in an inverse order, to determine in each profile the travel time function for the reflection plane corresponding to the just determined intersection point. The next stage is to allow variability in the fit, e.g. +/– half wavelength on each side of the predicted travel time function. The path of maximum signal energy/coherency is determined automatically within the interval allowed. For each interpreted reflector, a 3D reflector element is computed once the orientation of the reflector is completely determined. The stereographic representation method is not applicable to profiles measured in different boreholes because the intersection positions vary widely due to lateral extrapolation.

Figure 6-1. Example of reflector fitting procedure – reflector travel time curves for reflector 45, as interpreted from shot points 3, 8, 9 and 10 of KFM02A.

Easting

Figure 6-2. Wulff diagram for the interpretation of the orientation of reflector 45 determined from KFM02A.

Figure 6-3. Top view (left) and 3D view (right) of reflector elements of reflector 45, as determined from the VSP data acquired in KFM02A.

6.1 3D interpretation of reflectors from the VSP data acquired in borehole KFM01A

Figures 6-4 to 6-23 displays reflectors interpreted from the ten shot-points in borehole KFM01A. The results are discussed in chapter 7.

Figure 6-4. KFM01A. Maximum polarization component from SP01.

Figure 6-5. KFM01A. Reflectors interpreted from SP01.

Figure 6-6. KFM01A. Maximum polarization component from SP02.

Figure 6-7. KFM01A. Reflectors interpreted from SP02.

Figure 6-8. KFM01A. Maximum polarization component from SP03.

Figure 6-9. KFM01A. Reflectors interpreted from SP03.

Figure 6-10. KFM01A. Maximum polarization component from SP04.

Figure 6-11. Reflectors interpreted from SP04.

Figure 6-12. KFM01A. Maximum polarization component from SP05.

Figure 6-13. KFM01A. Reflectors interpreted from SP05.

Figure 6-14. KFM01A. Maximum polarization component from SP06.

Figure 6-15. KFM01A. Reflectors interpreted from SP06.

Figure 6-16. KFM01A. Maximum polarization component from SP07.

Figure 6-17. KFM01A. Reflectors interpreted from SP07.

Figure 6-18. KFM01A. Maximum polarization component from SP08.

Figure 6-19. KFM01A. Reflectors interpreted from SP08.

Figure 6-20. KFM01A. Maximum polarization component from SP09.

Figure 6-21. KFM01A. Reflectors interpreted from SP09.

Figure 6-22. KFM01A. Maximum polarization component from SP10.

Figure 6-23. KFM01A. Reflectors interpreted from SP10.

6.2 3D interpretation of reflectors from the VSP data acquired in borehole KFM02A

The Figures 6-24 to 6-41 display reflectors interpreted from ten shot-points in borehole KFM02A. The results are discussed in chapter 7.

Figure 6-24. KFM02A. Maximum polarization component from SP01.

Figure 6-25. KFM02A. Reflectors interpreted from SP01.

Figure 6-26. KFM02A. Maximum polarization component from SP02.

Figure 6-27. KFM02A. Reflectors interpreted from SP02.

Figure 6-28. KFM02A. Maximum polarization component from SP03.

Figure 6-29. KFM02A. Reflectors interpreted from SP03.

Figure 6-30. KFM02A. Maximum polarization component from SP04.

Figure 6-31. KFM02A. Reflectors interpreted from SP04.

Figure 6-32. KFM02A. Maximum polarization component from SP06.

Figure 6-33. KFM02A. Reflectors interpreted from SP06.

Figure 6-34. KFM02A. Maximum polarization component from SP07.

Figure 6-35. KFM02A. Reflectors interpreted from SP07.

Figure 6-36. KFM02A. Maximum polarization component from SP08.

Figure 6-37. KFM02A. Reflectors interpreted from SP08.

Figure 6-38. KFM02A. Maximum polarization component from SP09.

Figure 6-39. KFM02A. Reflectors interpreted from SP09.

Figure 6-40. KFM02A. Maximum polarization component from SP10.

Figure 6-41. KFM02A. Reflectors interpreted from SP10.

6.4 Nonconformities

The work has been executed according to the activity plan AP PF 400-04-60 in all aspects. There have been no nonconformities to note.

7 Results

The first column in Table 7-1 (for KFM01A) and Table 7-2 and Table B-2 (for KFM02A) is the event number, which is the same as the label of the reflector curves shown in Figure 6-5 to Figure 6-23 (for KFM01A) and Figure 6-24 to Figure 6-41 (for KFM02A). The numbering of the reflectors has been done independently for the two boreholes. The extension of the reflective elements shown in the 3-D figures is 100 m, which corresponds to $+/-$ two mean wavelengths.

The second column is the corresponding borehole length (or extension of the borehole length). This parameter is relevant for the interpretation of the reflectors that actually intersect the borehole. For the others, it is only a mathematical way of describing the position of the reflector relative to the borehole axis. The dips of the reflectors given in the third column and the dip directions in the fourth column were determined interactively, in several steps, as described in Chapter 6.

The last three columns give the coordinates of the crux point that, together with the coordinates of the origin chosen for interpretation (North = 6699000 , East = 1633000 ; Elevation $= 0$), fully characterizes the reflector. Having a common origin facilitates further integration of interpreted reflectors, among several boreholes and with those interpreted from surface seismic profiles. A Crux Point is defined as the foot of the perpendicular descended on a given reflection plane from an origin common for all profiles.

Appendix B lists reflector parameters for each of the boreholes, as interpreted from each shot point. In addition to columns 1 to 4 and 5 to 7 (which are equivalent to 1 to 4 and 7 to 9 respectively) in Table 7-1 (for KFM01A) and Table 7-2 (for KFM02A), Table B-1 and Table B-2 also provide the "Visibility" (column 5), "Event type" (column 6) and "Shot point" (column 10).

In each profile, reflectors are classified in three categories. Major events, appearing as well defined and continuous, belong to the first category (visibility mark $= 2 -$ marked with thick coloured line on the data profiles shown in Chapter 6). The weaker reflectors, visible but overridden by stronger events of other orientations belong to the second category (visibility mark $= 1 -$ marked with thin coloured line on the data profiles shown in Chapter 6). If the position and orientation of an event are determined from other profiles but the event does not appear as visible in the current profile, it is categorized as a third class (visibility mark $= 0$ – marked with interrupted coloured line on the data profiles shown in Chapter 6). The parameters of the visibility $= 0$ events are not given in Table B-1 and Table B-2. The mean value of the visibility marks, calculated from all profiles, is given for each reflector in Table B-1 and Table B-2, on the lines marked in bold face fonts. Reflectors with a mean mark larger than 0.8 (the absolute maximum being 1.0) are classified as strong (class I). Reflectors with a mean mark between 0.5 and 0.75 are classified as probable (class II). The weak seismic reflectors, with mean marks lower than 0.45 are classified as possible (class III).

Interpreted reflectors appear on the VSP profiles as direct P reflections or P-S conversions. This is documented by the "Event type" given in column 6 of Table B-1 and Table B-2.

7.1 Parameters of reflectors interpreted from KFM01A

The reflectors interpreted from KFM01A are summarized in Table 7-1. The dip is measured from the horizontal. To arrive at RT90 coordinates, add 6600000 to the Northing values and 1600000 to the Easting values.

Reflector label	Borehole length (m)	Dip (deg)	Dip direction (deg)	Reflector Class	Northing Crux	Easting Crux	Elevation Crux
$\mathbf{1}$	194	38	124	\mathbf{I}		32426	-900
$\overline{2}$	217	27	135	\mathbf{I}	99285	32712	-788
3	267	54	142	L	99822	32338	-765
4	324	56	181	L	99500	33009	-336
5	392	49	178	L	99536	32982	-474
6	465	49	165	L	99740	32791	-679
7	523	48	171	L	99701	32866	-637
8	572	24	161	L	99354	32879	-857
9	626	21	145	L	99317	32775	$-1,018$
10	751	23	135	L	99352	32638	$-1,195$
11	770	38	172	L	99674	32907	-859
12	832	34	161	L	99697	32759	$-1,079$
13	940	36	167	L	99756	32828	$-1,064$
14	1,050	37	163	L	99861	32743	$-1,180$
15	1,246	35	151	L	99876	32524	$-1,458$
16	1,317	27	152	L	99700	32623	$-1,594$
17	1,551	42	135	L	100093	31917	$-1,706$
18	1,590	35	166	L	100070	32732	$-1,552$
19	1,941	41	164	L	100392	32597	$-1,683$
20	2,024	38	159	L	100365	32472	$-1,888$
21	265	64	216	L	98740	32809	155
22	351	63	237	\mathbf{I}	98617	32409	363
23	675	70	234	L	98596	32450	245
24	2,803	82	238	\mathbf{I}	98548	32288	116
25	882	59	243	L	98756	32522	321
26	1,087	65	220	L	98912	32926	54
27	3,036	78	226	L	98828	32821	52
28	2,025	69	229	I	98892	32874	63
29	1,621	60	226		99043	33044	-35
30	1,977	61	223	I	99166	33157	-124
31	2,371	61	234	L	99085	33116	-80
32	2,953	64	216	L	99584	33423	-350
33	153	34	200	\mathbf{I}	99057	33007	-86
34	-18	32	187	Ш	99076	33003	-122
35	-1	32	194	Ш	99037	33009	-60
43	169	22	154	L	99199	32902	-551
45	311	35	172	L	99393	32937	-572
46	2,226	59	234	Ш	99100	33138	-104
49	913	61	207	Ш	99178	33090	-109

Table 7-1. VSP reflectors interpreted from borehole KFM01A.

As it can be observed from Figure 7-1, most gently dipping reflectors are dipping from N-NW to S-SE. Steeply dipping reflectors are observed in two groups, dipping from NE to SW and from N-NW to S-SE.

Figure 7-2 and 7-3 displays in a 3D view all reflector elements observed from borehole KFM01A.

Figure 7-1. Wulff polar (dip – azimuth) stereographic projection (see Figure A-3 and Figure A-4) of reflectors interpreted from all VSP profiles recorded from borehole KFM01A (blue dots) and their mean orientation (yellow dots). The vertical axis is North (Azimuth = 0°, Dip direction = 180°) – South and the horizontal is East (Azimuth = 90°, Dip direction = 270°) – West. The concentric black circles mark the locations of equal dip.

Figure 7-2. 3D view of all reflector elements interpreted from the VSP data acquired in borehole KFM01A. View from Southwest.

Figure 7-3. 3D view of all reflector elements interpreted from the VSP data acquired in borehole KFM01A. View from Northwest.

7.2 Parameters of reflectors interpreted from KFM02A

The reflectors interpreted from KFM02A are summarized in Table 7-2. The dip is measured from the horizontal. To arrive at RT90 coordinates, add 6600000 to the Northing values and 1600000 to the Easting values.

No	Depth (m)	Dip (deg)	Dip direction (deg)	Reflector Class	Northing Crux	Easting Crux	Elevation Crux
1	1,523	32	99	T	99100	32362	$-1,057$
2	1,296	42	130	I	99350	32584	-612
3	2,114	35	158	T	99856	32650	$-1,314$
4	1,390	55	181	Ш	99492	33010	-343
5	1,162	40	176	$\sf II$	99463	32970	-563
6	1,718	38	159	L	99697	32733	-942
7	1,737	42	163	I	99728	32779	-833
8	1,238	31	135	I	99338	32656	-797
9	1,431	33	123	T	99329	32501	-911
10	1,549	18	120	$\sf II$	99218	32622	$-1,338$
11	1,660	42	175	\mathbf{I}	99713	32933	-796
12	1,774	34	162	L	99711	32767	$-1,106$
13	1,811	31	159	L	99685	32739	$-1,229$
14	2,073	38	171	T	99909	32862	$-1,198$
15	2,211	34	149	I	99832	32501	$-1,422$
18	2,490	35	160	$\sf II$	100045	32617	$-1,574$
24	$-37,524$	84	237	$\ensuremath{\mathsf{II}}$	98532	32267	91
26	-73	65	220	Ш	98912	32926	54
27	428	88	264	Ш	98779	32750	8
28	-404	69	229	Ш	98890	32873	65
29	228	59	228	T	99032	33036	-29
30	727	59	217	\mathbf{I}	99175	33133	-133
31	422	61	234	$\ensuremath{\mathsf{II}}$	99084	33116	-79
32	2,857	70	217	$\ensuremath{\mathsf{II}}$	99594	33448	-265
33	185	34	198	T	99013	33007	-22
34	205	33	228	$\sf II$	99016	33005	-26
35	241	32	277	$\ensuremath{\mathsf{II}}$	99032	33006	-52
36	276	33	267	$\sf II$	99039	33000	-63
37	340	51	174	I	98989	33001	9
38	456	47	162	I	99059	32980	-60
39	529	25	137	I	99102	32902	-317
40	575	40	154	$\sf II$	99139	32933	-184
41	631	39	158	L	99172	32929	-230
42	699	37	125	I	99225	32963	-315
43	807	20	65	\mathbf{I}	99091	32986	-661
44	917	45	166	I	99309	32920	-320
45	982	38	171	L	99367	32941	-472
46	463	57	234	Ш	99103	33142	-114
47	672	62	74	$\ensuremath{\mathsf{II}}$	98934	32775	-124
48	595	63	80	$\sf II$	98968	32831	-90
49	808	62	207	Ш	99176	33091	-107

Table 7-2. VSP reflectors interpreted from borehole KFM02A.

As it can be observed from Figure 7-4, most gently dipping reflectors are dipping from N-NW to S-SE. Steeply dipping reflectors are observed in two groups, dipping from NE to SW and from N-NW to S-SE.

Figure 7-5 and 7-6 displays in a 3D view all reflector elements observed from borehole KFM02A.

Figure 7-4. Wulff polar (dip – azimuth) stereographic projection (see Figure A-3 and Figure A-4) of reflectors interpreted from all VSP profiles recorded from borehole KFM02A (dark blue dots) and their mean orientation (light blue dots). The vertical axis is North (Azimuth = 0°, Dip direction = 180°) – South and the horizontal is East (Azimuth = 90°, Dip direction = 270°) – West. The concentric black circles mark the locations of equal dip.

Figure 7-5. 3D view of all reflector elements interpreted from the VSP data acquired in borehole KFM02A. View from Southwest.

Figure 7-6. 3D view of all reflector elements interpreted from the VSP data acquired in borehole KFM02A. View from Northwest.

7.3 Combination of reflectors from KFM01A and KFM02A

All reflector elements described in Section 7.1 and Section 7.2 from the two boreholes KFM01A and KFM02A are merged together in a 3D view in the Figures 7-7 and 7-8, as a top view in Figure 7-9 and in a Wulf stereographic projection in Figure 7-10.

Figure 7-7. 3D view of all reflector elements interpreted from the VSP data acquired in boreholes KFM01A and KFM02A. View from Southwest.

Figure 7-8. 3D view of all reflector elements interpreted from the VSP data acquired in boreholes KFM01A and KFM02A. View from Southeast.

Figure 7-9. Top view of all reflector elements interpreted from the VSP data acquired in boreholes KFM01A and KFM02A.

Figure 7-10. Wulff polar (dip – azimuth) stereographic projection of reflectors interpreted from all VSP profiles recorded from boreholes KFM01A and KFM02A (dark blue dots) and their mean orientation (yellow for KFM01A and light blue dots for KFM02A). The vertical axis is North (Azimuth = 0°, Dip direction = 180°) – South and the horizontal is East (Azimuth = 90°, Dip direction = 270°) – West. The concentric black circles mark the locations of equal dip.

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Appendix A

A.1 The Image Point (IP) transform

The IP (Image Point) forms the core of the technique that we have developed for filtering crystalline rock VSP and cross-hole data. The Radon transforms have been applied to seismic data in many ways, e.g. slant stack τ*-p* transform, where the apparent velocity limits are varied according to the predicted travel times of reflections. Unlike with the τ*-p*, with the IP transform the stacking is done along hyperbolic paths corresponding to the timedepth functions of possible reflectors. Due to this "natural" choice of the stacking paths, the coherency can be used effectively to enhance the weak reflections. Other advantages of the IP transform include the efficient separation of interfering reflections and the possibility of using directly the IP transformed profiles as interpretation tools.

In a constant velocity medium, a reflection plane is completely defined by the mirror image of the source S with respect to that plane, i.e. the Image Point, as shown in Figure A-1. Therefore, one can generate any time-depth function associated with a planar reflector by computing the distance from the Image Point to each detector and divide it by the velocity. Therefore, the IP transform turns the hyperbolic reflection patterns of the time-depth space into points in the IP space.

A cylindrical co-ordinate system (ξ , ζ , φ) has been used in Figure A-1, where the ζ -axis lies along the borehole (assumed to be straight), the ξ-axis is perpendicular on the ζ-axis and contains the source S and the angle φ is the azimuth relative to the ζ -axis. All image sources with the same (ξ, ζ) co-ordinates produce identical time-depth functions because the distances from the image source to all receivers are independent of the relative azimuth φ . This property allows most of the computations to be performed in two dimensions. The third dimension is regained at a later stage, when the polarisation of the data is examined.

Figure A-1. Cylindrical co-ordinates system used in the Image Point space.

The time-depth function of a reflected event in a time-depth profile is

$$
t_r = \sqrt{(\xi^2 + (z - \zeta)^2)} / \nu = \sqrt{\xi^2 + \zeta^2 + z^2 - 2z\zeta} / \nu
$$
 (A.1)

where t_r is the travel time to the receiver placed at depth z , v is the constant propagation velocity and (ξ, ζ) are the co-ordinates of the Image Source.

Equation (A.1) becomes

$$
t_r = \sqrt{\rho^2 + z^2 - 2z\zeta} / v \tag{A.2}
$$

by replacing the variable ξ with

$$
\rho = \sqrt{\xi^2 + \zeta^2} \tag{A.3}
$$

If the distance ρ from the origin is used instead of the offset ξ , the transformed representation will not be stretched and easier to use as an interpretive tool. The travel time does not depend on the relative azimuth φ , i.e. the locus of possible image points forms a circle centred on the borehole.

The Image Point transformed profile $\Gamma(\rho, \zeta)$ is obtained by stacking in the time-depth profile $g(z, t)$ along paths corresponding to all (ρ, ζ) pairs within a selected domain $\rho_{\min} \le \rho \le \rho_{\max}$, $\zeta_{\min} \leq \zeta \leq \zeta_{\max}$, thus covering all possible positions and orientations that a reflecting plane may have:

$$
\Gamma(\zeta,\rho) = \int_{Z_{\min}}^{Z_{\max}} dz \ g(z,t=t_r(\zeta,\rho;z)) \tag{A.4}
$$

where $t_r(\zeta, \rho; z)$ is the arrival time corresponding to the planar reflector specified by ρ and ζ , to the detector at depth *z*.

The inverse transform is constructed by integrating along each path that received the contribution of $g(z, t)$ in the direct transform:

$$
g(z,t) = \frac{1}{2\pi \cdot v^2 \cdot t} H \frac{\partial}{\partial t} \int_{\zeta_1}^{\zeta_2} d\zeta \Gamma(\zeta; \rho = \rho_r(z,t;\zeta))
$$
 (A.5)

where

$$
\rho_r = \sqrt{v^2 t^2 - z^2 + 2z\zeta}
$$
 (A.6)

The Hilbert transform **H** and the derivation with respect to time are used to restore the original signal shape, similarly as with the τ*-p* transform.

The inverse transform from Image Space back to depth-time space leads to a filtered version of the reflection profile. This filtering effect cannot be avoided, as the summing is done only along paths corresponding to time functions of true reflected signals travelling with the velocity *v*. These events stack properly while noise and other wave types are suppressed.

A.2 Polarisation analysis in the Image Point space

The (ρ, ζ) co-ordinates of the Image Point space define two of the three parameters needed to determine the 3-D position of a reflector (the azimuth φ relative to the ζ -axis plays no role in the IP transform).

A straightforward method to estimate the azimuth is to rotate incrementally the **R** and **T** components and observe at which angle reflection patterns disappear. This angle would give the perpendicular direction to the actual signal polarization. A more elaborate technique relies on solving the eigenvector problem to determine the direction and the degree of polarization in a time window. A filtering effect is achieved by projecting the instantaneous particle motion at the time t on the average polarization direction determined in a time window centered on t. Multiplying the projected components, with $L(t) = I - (\lambda_2(t)/\lambda_1(t))^2$, where $\lambda_i(t)$ and $\lambda_2(t)$ are the largest and the second largest eigenvalues determined in the respective window, would enhance the polarized events.

A problem encountered when attempting to apply polarization analysis to crystalline rock VSP data is the intermingling and criss-crossing of events arriving from diverse directions. In fact, there would be very few, if any, time windows in a time-depth profile where a reflected arrival is unaffected by interference with other arrivals and noise. The result is that the polarization estimated in the time-depth space is, generally, extremely unstable. Polarization analysis performed in the Image Point space largely avoids such problems because the hyperbolic patterns from the time-depth profiles collapse to separate vicinities in the IP space.

A.3 Non-Linear enhancement of the reflectors

The general filtering effect of the Image Point transform results from the use of the actual propagation velocity in the computation of the integral paths. The filtering effect can be increased by following a non-linear approach instead of the linear integral in the inverse transform. For example, the amplitudes along each stacking path can be ordered and the sum be performed only for a chosen subset. A less computationally demanding method is to compute the envelopes in the IP transformed space and weight the sum by the maximum of the envelope along each integration path. The optimum weight to be used with the filter depends on certain characteristics of the data and is to be decided upon on a case-by-case basis. Large weights of the envelope maxima would result in well defined but fewer events, the fainter reflections being filtered out along with the noise. Small weights would let through more events but the coherency would improve only slightly. A fairly safe procedure for preventing over-processing consists of randomly rearranging bits and parts of the IP transformed profile, into a "dummy" one, so that causal coherency is destroyed. The dummy profile is then run through the same processing scheme as the real profile. Coherent events seeming to emerge in the dummy profile indicate that the power of the filter must be turned down.

A.4 Multi-profile fitting

The azimuth estimates obtained by polarization analysis are improved by the concurrent processing of several profiles. VSP investigations are normally conducted in several boreholes, each layout comprising typically ten offsets. To facilitate the consistent interpretation of the considerable number of reflection events, an attempt is made to link the ones likely to represent the same feature in different profiles. However, this reduction procedure cannot easily be done only interactively, as the gap between layouts is often large and the site features are not necessarily planar. Solving the problem requires the use of statistical methods, e.g. clustering analysis.

If the profiles are measured in the same borehole, the mean orientation of a reflector can be found in a stereographic representation (Wulff diagram) at the intersection of the dip-azimuth curves computed for each profile, as shown in Figure A-2. As reflectors associated with fracture zones are generally nearly planar, they are assumed to intersect the

borehole (or its extension) at the same depth. The stereographic representation method is not applicable to profiles measured in different boreholes because the intersection positions vary widely due to lateral extrapolation. The 3D fitting procedure exemplified in Figure A-5 is applied instead. The curves in Figure A-6 are described by the foot of the perpendicular descended on a given reflection plane from chosen origin, common for all profiles.

Figure A-2. Different shot gathers filtered by the IP transform.

Figure A-3. The Wulff equal-angle stereographic projection of a reflector is given by the intersection of the segment drawn between the North Pole (P) and the point on the lower hemisphere, which has the coordinates equal to the dip and azimuth of the reflector element with the Equatorial plane. Azimuth = Dip direction +180°.

Figure A-4. Different stereographic projections: Wulff (the two on the left) equal-angle, polar and equatorial and Schmidt/Lambert (the two on the right) equal-area, polar and equatorial.

Figure A-5. Wulff diagram representing the same reflector as the one depicted in Figure A-2.

Figure A-6. Crux point diagram representing the same reflector as the one depicted in Figure A-2.

Detailed description of all reflectors observed from KFM01A (Table B-1) and from KFM02A (Table B-2). The dip is measured from the horizontal. To arrive at RT90 coordinates, add 6600000 to the Northing values and 1600000 to the Easting values.

No	Borehole length (m)	Dip $(°)$	Dip dir $(°)$	Visibility marks	Event type	Delta depth (m)	Delta Dip (°)	Delta Dip dir (°)	Shot point
$\mathbf{1}$	194.3	37.8	118.9	$\mathbf{1}$	P reflection	43.4	25.9	35.0	KFM01A-01
1	194.3	37.8	118.9	\overline{c}	PS conversion	15.1	20.0	161.5	KFM01A-02
1	194.3	37.8	118.9	$\overline{2}$	P reflection	13.6	18.7	33.7	KFM01A-03
1	194.3	37.8	118.9	$\overline{2}$	P reflection	17.7	10.0	67.4	KFM01A-04
1	194.4	37.8	148.9	$\overline{2}$	P reflection	13.9	1.3	25.7	KFM01A-05
1	194.5	37.8	108.9	\overline{c}	P reflection	21.6	20.6	57.1	KFM01A-06
1	194.3	37.8	118.9	\overline{c}	P reflection	20.5	16.2	45.0	KFM01A-07
1	194.4	37.8	138.9	$\overline{2}$	P reflection	21.5	16.5	44.9	KFM01A-08
1	194.3	37.8	118.9	$\mathbf{1}$	P reflection	86.4	37.8	61.1	KFM01A-09
1	194.3	37.8	118.9	$\mathbf{1}$	P reflection	113.4	46.7	91.0	KFM01A-10
1	194.3	37.8	122.9	0.85		13.6	1.3	25.7	
2	217.2	25.0	134.4	$\overline{2}$	P reflection	25.6	17.1	85.0	KFM01A-02
\overline{c}	217.2	25.0	134.4	$\overline{2}$	P reflection	11.4	16.0	43.6	KFM01A-03
2	217.2	31.0	134.4	$\overline{2}$	P reflection	13.5	1.1	50.8	KFM01A-04
\overline{c}	217.2	29.6	144.4	$\overline{2}$	P reflection	13.4	6.5	21.1	KFM01A-05
$\overline{\mathbf{c}}$	217.4	25.0	125.0	\overline{c}	P reflection	18.9	29.3	39.7	KFM01A-06
$\overline{2}$	217.3	25.0	135.0	$\overline{2}$	P reflection	19.8	30.7	29.8	KFM01A-07
\overline{c}	217.2	29.6	134.4	$\overline{2}$	P reflection	22.7	30.7	49.5	KFM01A-08
$\mathbf{2}$	217.2	27.2	134.6	0.7		11.4	1.1	21.1	
3	267.0	53.0	135.0	\overline{c}	PS conversion	24.6	11.3	78.1	KFM01A-02
3	267.0	53.0	135.0	$\overline{2}$	PS conversion	11.8	36.4	14.4	KFM01A-03
3	267.0	53.0	135.0	$\overline{2}$	P reflection	19.7	1.7	49.9	KFM01A-04
3	267.0	55.7	147.4	$\overline{2}$	P reflection	15.8	6.4	2.6	KFM01A-05
3	267.6	55.7	137.4	\overline{c}	P reflection	16.6	6.0	25.9	KFM01A-06
3	266.9	55.7	157.4	$\overline{2}$	P reflection	16.3	13.3	0.7	KFM01A-07
3	267.0	53.0	150.0	$\overline{2}$	P reflection	18.9	1.9	33.3	KFM01A-08
3	267.0	55.7	137.4	$\mathbf{1}$	PS conversion	398.4	55.7	42.6	KFM01A-09
3	267.0	55.7	137.4	\overline{c}	PS conversion	18.9	18.2	85.3	KFM01A-10
3	267.0	54.5	141.3	0.9		11.8	1.7	0.7	
4	323.9	57.0	183.2	2	PS conversion	18.1	21.0	24.4	KFM01A-02
4	323.9	57.0	183.2	2	PS conversion	17.8	18.5	15.1	KFM01A-03
4	323.9	57.0	183.2	2	P reflection	18.8	3.9	0.4	KFM01A-04
4	323.9	57.0	183.2	2	P reflection	25.4	9.6	29.5	KFM01A-05
4	324.4	57.0	173.2	2	P reflection	16.7	3.5	7.9	KFM01A-06
4	323.9	57.0	183.2	$\overline{2}$	P reflection	18.4	0.8	17.1	KFM01A-07
4	324.1	57.0	180.0	2	P reflection	17.0	5.9	2.8	KFM01A-08
4	324.2	50.0	180.0	2	P reflection	24.6	50.0	0.0	KFM01A-09

Table B-1. Geometrical parameters of reflectors interpreted from VSP data acquired in borehole KFM01A. The lines written in bold fonts give the mean values for each parameter, as computed for each reflector.

		49 913.2 60.0 206.9 1 Preflection 20.4 1.5 7.6 KFM01A-10		
	49 913.2 60.8 206.9 0.1		20.4 1.5 7.6	

Table B-2. Geometrical parameters of reflectors interpreted from VSP data acquired in borehole KFM02A. The lines written in bold fonts give the mean values for each parameter, as computed for each reflector.

